Evolution of a dense outer protective tube layer in serpulids (Polychaeta, Annelida)

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Abstract: Although the walls of most serpulid tubes are homogeneous, tubes of certain species may contain up to four ultrastructurally distinct layers. Some of these layers are made of densely packed large crystals and others are composed of sparsely packed fine crystals. In almost all (16 of 17) examined species having layered tubes, the dense layer is located in the outer wall part and the layer(s) composed of fine and relatively sparsely packed crystals are positioned in the inner wall part. Two species have transparent tube walls made entirely of densely packed crystals. Fossil serpulid tubes with dense outer layers (DOL) are known from the Late Cretaceous (*Pentaditrupa subtorquata*) and the Eocene (*Pyrgopolon* cf. *mellevillei* and *Rotularia spirulaea*). DOL gives a characteristic smooth shiny appearance to the tube surface and presumably evolved as an adaptation against drilling predation by gastropods and to delay shell dissolution in the waters of the deep-sea under-saturated with calcium carbonate.

Key Words: Serpulidae; biomineralization; tube ultrastructure; evolution.

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Résumé : Évolution de la couche externe, dense et protectrice, du tube de Serpulidés (Polychètes, Annélides).- Alors que les parois de la plupart des tubes de Serpulidés sont homogènes, les tubes de quelques espèces peuvent présenter jusqu'à quatre couches à ultrastructures variées. Quelques-unes de ces couches sont constituées de gros cristaux compacts et d'autres de fins cristaux épars. Chez presque toutes les espèces étudiées ayant un tube stratifié (16 sur 17), la couche compacte se situe dans la partie externe de la paroi minéralisée du tube ; quant à celle(s) constituée(s) de cristaux fins et présentant un agencement relativement lâche, elle(s) se trouve(nt) dans la partie interne de la paroi. Deux espèces présentent des parois transparentes constituées de cristaux compacts. Les tubes de Serpulidés fossiles à couches externes compactes ("dense outer layers", DOL) sont connus dès le Crétacé supérieur avec *Pentaditrupa subtorquata*, et à l'Éocène avec, par exemple, *Pyrgopolon* cf. *mellevillei* et *Rotularia spirulaea*. Cette ultrastructure, DOL, donne un aspect lisse et brillant, caractéristique, à la surface du tube et pour retarder la dissolution de la paroi minéralisée dans des eaux sous-saturées en carbonate de calcium.

Mots-Clefs : Serpulidae ; biominéralisation ; ultrastructure du tube ; évolution.

Introduction

Serpulids are polychaetes that build elaborate calcareous tubes. This large and highly successful group has a cosmopolitan geographic and wide bathymetric distribution, from the intertidal to the abyssal zones (ten Hove & KUPRIYANOVA, 2009; KUPRIYANOVA *et alii*, 2011, KUPRIYANOVA & NISHI, 2011) and its representatives are even found in chemosynthetic communities of hydrocarbon seeps and hydrothermal vents (KUPRIYANOVA *et alii*, 2010). Some taxa (of the genus *Ficopomatus*) tolerate a wide range of salinities (ten Hove & KUPRIYANOVA, 2009) and one species, *Marifugia cavatica*, inhabits the fresh waters of subterranean caves (KUPRIYANO-VA *et alii*, 2009).

Serpulid tubes are composed of either calcite (most stable common natural polymorph of CaCO₃), aragonite (another common polymorph of CaCO₃), or a combination of both (LOWEN-STAM, 1954; BORNHOLD & MILLIMAN, 1973; SIMKISS & WILBUR, 1989; VOVELLE *et alii*, 1991; VINN *et alii*, 2008a, b, c). Because of their calcareous tubes, serpulids may have played an important role in the ocean carbonate sink (both ongoing and past) as their tubes contribute to the sedimentary rock formation (MASTRANGELO & PASSERI,

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Species	Clade	Multilayered tube with densely and sparsely packed crystals/ type of		Dense outer layer present or whole tube made of dense		Material studied
		structures		structure/type of structure		
Apomatus globifer	в	No	IOP	No		1
Bathyvermilia langerhansi	В	Yes	HAC/HAC	Yes	HAC	2
Chitinopoma serrula	В	No	IOP	No		3
Chitinopomoides wilsoni	В	No	SIOP	No		4
Crucigera websteri	A	Yes	SPHP/IOP/LF/SPHP	Yes	SPHP	5
Crucigera zygophora	A	Yes	SPHP/IOP/SIOP	Yes		6
Ditrupa arietina	А	Yes	RRP/IOP	Yes	RRP	7
Ditrupa gracillima	A	Yes	RRP/IOP	Yes	RRP	8
Ditrupa strangulata	A	Yes	RRP/FH	Yes	RRP	9
Ficopomatus enigmaticus	A	No	IOP	No		10
Ficopomatus uschakovi	A	No	IOP	No		11
Filograna implexa	В	No	IOP	No		12
Filogranella elatensis	В	No	IOP	No		13
Filogranula gracilis	В	No	IOP	No		14
Floriprotis sabiuraensis	A	No	IOP/LF/SLF	No		15
Galeolaria hystrix	A	No	LF/SIOP	No		16
Hyalopomatus marenzelleri	A	No	IOP	No		17
Hyalopomatus madreporae	A	NO	IOP	No		18
Hydroides dianthus	A	Yes	SPHP/IOP/LF	Yes	SPHP	19
Hydroides norvegicus	A	Yes	SPHP/IOP	Yes	SPHP	20
Hydroides spongicola	A	No	SIOP	No		21
Janita fimbriata	В	No	IOP	No		22
Josephella marenzelleri	в	NO	IOP	NO		23
Laminatubus alvini	A	Yes	HAC/IOP	Yes	HAC	24
Matnugia cavalica	A B	No	IOP/IOP	No		20
Neovermilia falsigera	A2	Vec		Vec	TOPI	20
Neovermilia	A: 	No	I F	No	IOFL	27
sphaeropomatus		No	TOP	No		20
Paraprotis pulchra	В	NO		NO Vez	CDUD	29
Pentaditrupa subtorquata	r A	Tes	SPRP/FR	Yes	SPHP	30
Pomatostegus stellatus	B	No	HAC	No	- SF	32
Protis arctica	В	No	TOP	No		33
Protula diomedeae	В	No	SOLOP	No		34
Pseudovermilia madracicola	В	No	SIOP	No		35
Pseudovermilia occidentalis	в	No	IOP	No		36
Pyrgopolon ctenactis	А	No	SOSIOP	No		37
Rhodopsis pusilla	В	No	IOP	No		38
Rotularia spirulaea	?	Yes	HAC/LF	Yes	HAC	39
Salmacina incrustans	В	No	IOP	No		40
Semivermilia crenata	В	No	IOP	No		41
Serpula crenata	А	Yes	SP/IOP	Yes	SP	42
Serpula israelitica	A	No	LF	No		43
Serpula vermicularis	A	No	LF	No		44
Spiraserpula caribensis	А	Yes	SPHP/SIOP/SPHP	Yes	SPHP	45
Spirobranchus giganteus	A	No	OF/SIOP	No		46
Spirobranchus triqueter	A	No	LF	No		47
Spirobranchus kraussii	A	Yes	SPHP/IOP/LF/SIOP	Yes	SPHP	48
Vermiliopsis infundibulum	В	No	SIOP	No		49
Vitreotubus digeronimoi	A	No**	SP	Yes	SP	50

** Single layered species with tube walls made entirely of densely packed crystals.

Table 1: Serpulid species studied in this paper and their tube ultrastructures.

No Yes: absence or presence of "Multilayered tube with densely and sparsely packed crystals".

Isotropic structures: HAC - homogeneous angular crystal structure, FH - fine grained homogeneous structure, IOP – irregularly oriented prismatic structure, IOPL - irregularly oriented platy structure, RHC - rounded homogeneous structure, SIOP - spherulitic irregularly oriented prismatic structure. Semi-oriented structures: SOIOP - semi-ordered irregularly oriented prismatic structure, SOSIOP - semi-ordered spherulitic oriented prismatic structure. Oriented prismatic structures: RRP- regularly ridged prismatic structure, SP - simple prismatic structure, SPHP – spherulitic prismatic structure.

Oriented complex structures: LF -lamello-fibrillar structure, OF - ordered fibrillar structure, SLF - spherulitic lamello-fibrillar structure. Tube layers are ordered from outside (left) to lumen (right).

1. Apomatus globifer. Recent. Locality: Kara Sea, 71°N, 64°E, depth 122 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.0105.01)

2. Bathyvermilia langerhansi. Recent. Locality: Madeira, Porto Santo, St.4.180, depth 3499 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4041)

3. *Chitinopoma serrula*. Recent. Locality: Iceland, Snaefellsnes Peninsula, depth 30 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.5034)

4. *Chitinopomoides wilsoni*. Recent. Locality: Antarctica, USNM Acq.224443, depth 80 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3166)

5. *Crucigera websteri*. Recent. Locality: Surinam, depth 60 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3589)

6. Crucigera zygophora. Recent. Locality: Canoe Bay, Alaska, USA, depth 8 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3287)

7. Ditrupa arietina. Recent. Locality: Sweden, Tjärnö, depth 10 m. Deposited: Natural History Museum, University of Tartu

8. *Ditrupa gracillima*. Fossil. Locality: Australia, VIC, Rutledge Creek, Middle Miocene (Bairnsdalian) Rutledge Creek 9. Member, presumable palaeo water depth 10 m. Deposited: Zoological Museum Amsterdam.

9. *Ditrupa strangulata*. Fossil. Locality: France, Paris Basin, Belleu near Soissons, Eocene (Middle Lutetian). Deposited: Zoological Museum Amsterdam (ZMA V.Pol. 3115)

10. *Ficopomatus enigmaticus*. Recent. Locality: Lake of Tunis, depth 8 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3779)

11. *Ficopomatus uschakovi*. Recent. Locality: Thailand, Kong Prao, Koh-Chang Island, depth 1-2 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3647)

12. *Filograna implexa*. Recent. Locality: UK, Orkney Island, depth 22 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3767)

13. *Filogranella elatensis*. Recent. Locality: Japan, Sesoko Island, Okinawa, depth 10 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3911)

14. *Filogranula gracilis*. Recent. Cape Verde Islands, Boa Vista, depth 111 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4047)

15. *Floriprotis sabiuraensis*. Recent. Locality: Shimoshima Island, Amakusa, Japan, depth 10 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3929)

16. Galeolaria hystrix. Recent. Locality: New Zealand, Queen Charlotte Sound, depth 1-2 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3576)

17. Hyalopomatus marenzelleri. Recent. Locality: Canary Islands, Lanzarote, depth 1030-1070 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4522)

18. Hyalopomatus madreporae. Recent. Locality: Santa Maria di Leuca, Italy, depth 497-790 m.

19. Hydroides dianthus. Recent. Locality: USA, Anna Maria Island, FL, depth 2 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3661)

20. *Hydroides norvegicus*. Recent. Norway, Skjerstad (Saltenfjord), depth 164 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.0463).

21. *Hydroides spongicola*. Recent. Locality: Netherlands Antilles, Curaçao, depth 7 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3584)

22. Janita fimbriata. Recent. Locality: Canary Islands, Lanzarote, depth 88 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4072)

23. *Josephella marenzelleri*. Recent. Locality: France, Marseille, depth unknown. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3030)

24. Laminatubus alvini. Recent. Locality: East Pacific Rise, 09°N, 104°W, depth 2509 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3512)

25. *Marifugia cavatica*. Recent. Locality: Hercegovina, Popovo Polje fresh water caves, depth 1 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3102)

26. *Metavermilia multicristata.* Recent. Locality: Seychelles, N. of d'Arros Island, depth 55 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4300)

27. *Neovermilia falcigera*. Fossil. Locality: Italy, Calabria, Pleistocene, depth 1580 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3501b)

28. *Neovermilia sphaeropomatus*. Recent. Locality: New Zealand, Cape Saunders, depth 10 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3274)

29. *Paraprotis pulchra*. Recent. Locality: Japan, Kushimoto, depth 70 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3379)

30. *Pentaditrupa subtorquata*. Fossil. Locality: Upper Campanian, Cretaceous of Höver by Hannover, Germany. Deposited: Zoological Museum Amsterdam (ZMA V.Pol. 3705)

31. *Placostegus tridentatus.* Recent. Locality: Norway, Bergensfjord, depth unknown. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.1105)

32. Pomatostegus stellatus. Recent. Locality: Netherlands Antilles, Curaçao, depth 0.5 m. Deposited: Zoological Mu-

seum Amsterdam (ZMA V.Pol.5170)

33. Protis arctica. Recent. Locality: NE of Iceland, depth 1802 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3833)

34. Protula diomedeae. Recent. Locality: USA, Florida, depth 73 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4271)

35. *Pseudovermilia madracicola*. Recent. Locality: Netherlands Antilles, Curaçao, Salinja Fuik, depth 27 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3751)

36. *Pseudovermilia occidentalis*. Recent. Locality: Cape Verde Islands, St. Luzia, depth 10 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4090)

37. *Pyrgopolon ctenactis*. Recent. Locality: Netherlands Antilles, Bonaire, depth 15 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4969)

38. *Rhodopsis pusilla*. Recent. Locality: Reunion Island, depth 5 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3623)

39. *Rotularia spirulaea*. Fossil. Locality: Eocene, Doss Trento, Italy. Deposited: Natural History Museum Vienna (NHMW 2006z0233/0013)

40. *Salmacina incrustans*. Recent. Locality: Spain Costa Brava, depth 0.5 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3814)

41. Semivermilia crenata. Recent. Locality: France, Marseille, depth 1 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol. 3045)

42. *Serpula crenata*. Recent. Locality: Indonesia, 411 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.1739)

43. Serpula israelitica. Recent. Locality: Cape Verde Islands, São Vicente, depth unknown. Deposited: Zoological Museum Amsterdam.

44. *Serpula vermicularis*. Recent. Locality: Ireland, Ardbear Lough, depth 20 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.3780)

45. Spiraserpula caribensis. Recent. Locality: Netherlands Antilles, Curaçao, depth 1 m. Deposited: Zoological Museum Amsterdam

46. Spirobranchus giganteus. Recent. Locality: Netherlands Antilles, Curaçao, depth 6 m. Deposited: Zoological Museum Amsterdam

47. Spirobranchus triqueter. Recent. Locality: Sweden, Tjärnö, depth 10 m. Deposited: Natural History Museum, University of Tartu.

48. Spirobranchus kraussii. Recent. Locality: Teluk Slawi, Indonesia, depth 0.5 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4748)

49. Vermiliopsis infundibulum. Recent. Locality: Canary Islands, Lanzarote, depth 80 m. Deposited: Zoological Museum Amsterdam (ZMA V.Pol.4036)

50. Vitreotubus digeronimoi. Recent. Locality: Seychelles, Platte Island, Sta.795, depth 600 m. Zoological Museum Amsterdam (ZMA V.Pol.4308)

1975; ten Hove & Van den Hurk, 1993; MEDER-NACH *et alii*, 2000; VINN & MUTVEI, 2009). The group has the best fossil record among all annelids. The earliest undisputable records of serpulids belong to the Middle Triassic, while earlier Palaeozoic records of the serpulids belong to various groups of problematic tubeworms (BURCHETTE & RIDING, 1977; WEEDON , 1994; VINN & MUTVEI, 2005, 2009; VINN, 2006, 2010; VINN & ISAKAR, 2007; VINN & MÖTUS, 2008).

The tubes of serpulids have a relatively simple external macromorphology. Internally, the mineral tubes are lined with a thin organic layer (NISHI, 1993; VINN, 2011), but a thick external organic (conchiolin) layer, such as found in the periostracum of molluscs and brachiopods is absent. On the inside, the tube surface is usually smooth, although some species form internal transverse tabulae (HEDLEY, 1958; ten Hove, 1973; ten Hove & Smith, 1990), or rarely, longitudinal keels (Spiraserpula: PILLAI & ten Hove, 1994). Growth lamellae, the units of tube wall produced during a single secretion episode, are commonly chevron shaped in longitudinal section (WEEDON, 1994; VINN & FUR-RER, 2008) or, more rarely, are straight bands (SANFILIPPO, 1996). At the point of contact between the lateral tube wall and the substrate, the tube may produce cavities known as alveolar structures (e.g., HEDLEY, 1958).

Externally, serpulid tubes may bear additional sculpturing in the form of smooth or spiny longitudinal keels and transverse ridges or peristomes (REGENHARDT, 1964; ten Hove, 1973, 1975; BASTIDA-ZAVALA & ten HOVE, 2002; ALIANI *et alii*, 1995; ten Hove & KUPRIYANOVA, 2009). The surface of most tubes shows faint growth striations, however, in a number of species it is smooth and appears to be covered with a transparent layer that gives the tube porcelain appearance (IPPOLITOV & RZHAVSKY, 2008; ten HOVE & KUPRIYANOVA, 2009).

In contrast to relatively simple morphology, Recent and fossil serpulid tubes may show surprising ultrastructural complexity (VINN, 2007; VINN & FURRER, 2008; VINN *et alii*, 2008a, b, c, d): at least thirteen distinct tube ultrastructures involving unoriented, uniformly oriented, or complex oriented crystals have been described (VINN *et alii*, 2008b, d). Although most serpulid tubes are single-layered, there can be up to four ultrastructurally distinct layers (VINN *et alii*, 2008b).

There are two controversial models of biomineralization in serpulids. According to the traditional model (SIMKISS & WILBUR, 1989), the calcareous granules excreted as slurry from the glands underneath the collar solidify slowly, allowing the folded collar to mould the calcitesaturated mucus into the end tube part. However, because this model cannot explain the observed ultrastructural complexity of the serpulid tubes, recent studies of tube ultrastructure (VINN *et alii*, 2008b, 2009) and biochemistry (TANUR *et alii*, 2010) resulted in new models that explain biomineralization as a organic matrix-mediated process.



Figure 1: *Pyrgopolon mellevillei* (NHMW 2006z0233/ 0005) London Clay, Eocene, England. Polished cross section of the tube, treated with the MUTVEI solution for 5 minutes, showing outer dense coarse crystal layer. SPHP - spherulitic prismatic structure; IOP – irregularly oriented prismatic structure.



Figure 2: *Pyrgopolon mellevillei* (NHMW 2006z0233/ 0005) London Clay, Eocene, England. Polished cross section of the tube through the outer dense tube layer (spherulitic prismatic structure), treated with the MUT-VEI solution for 5 minutes. Ext - exterior.



Figure 3: *Pyrgopolon mellevillei* (NHMW 2006z0233/ 0005) London Clay, Eocene, England. Polished cross section of the tube through the inner sparsely packed fine crystal tube layer (irregularly oriented prismatic structure), treated with the MUTVEI solution for 5 minutes.

The aim of the study was to examine whether there is a pattern in the position of various tube layers. We also discuss functions of the outer tube layer and potential trends in evolution of dense layers in serpulid tubes.

Material and methods

This study is based on an extensive database containing scanning electron micrographs (SEM) of tube ultrastructures of 45 globally distributed Recent and 20 fossil serpulid species ranging in age from the Lower Jurassic through the Pliocene of Europe (SANFILIPPO, 2009; VINN, 2005, 2007; VINN et alii, 2008b, c, d). The database was examined to determine the structural differences between the outer and inner parts of the serpulid tube wall (Table 1). Recent material was collected intertidally or by diving, trawling, or dredging; it was fixed in 4% (buffered) formalin and was later transferred to 70% ethanol for museum storage. The material was deposited in Zoological Museum, University of Amsterdam, the Netherlands (ZMA), now in the process of being merged with the Netherlands Centre for Biodiversity, Naturalis, Leyden. Serpulid species were selected for the study in order to cover most of the accepted genera.

All samples with a mineralized ultrastructure (= tube fragments) were embedded in epoxy resin, polished in longitudinal and transversal directions, and then etched with 1% acetic acid for five to ten minutes prior to scanning electron microscope (SEM) examination. We used 5-15-mm-long longitudinal sections, and between one and three cross sections of each serpulid species. Some samples were repolished and treated with a 1:1 mixture of 25% glutaraldehyde and 1% acetic acid, to which alcian blue was added (MUTVEI solution) before performing the SEM study (SCHÖNE et alii, 2005). As a result of this treatment, the organic-rich parts of the tubes have a more intensive blue color. Some samples were repolished and bleached with NaCIO for removing organic matrixes before the SEM study. Scanning electron microscopy was done with a Hitachi S-4300, equipped with an Inca EDX system (energy-dispersive X-ray spectroscope), at the Swedish Museum of Natural History, Stockholm, and the etched samples and organic lining were viewed with a Zeiss 940D microscope, equipped with SAMx SDD EDX (energy-dispersive X-ray spectroscope), at the Department of Geology, University of Tartu. The beam was operated at 5-10 kV and 1 nA.

Evolution of serpulid shell layers was also analyzed using a published cladogram of serpulid relationships (KUPRIYANOVA & NISHI, 2010). This phylogenetic tree was selected for the present study because it covers the best range of genera of this paper.



Figure 4: Laminatubus alvini, East Pacific Rise, Recent. Polished cross section of the tube, showing outer dense coarse crystal layer (HAC) and inner sparsely packed fine crystal layer (IOP). HAC – homogeneous angular crystal structure; IOP - irregularly oriented prismatic structure.



Figure 5: *Neovermilia falcigera*, Italy, Calabria, Pleistocene. Polished cross section of the tube through outer dense coarse crystal tube layer (irregularly oriented platy structure), treated with MUTVEI solution for 5 minutes.



Figure 6: *Neovermilia falcigera*, Italy, Calabria, Pleistocene. Polished cross section of the tube through inner sparsely packed fine crystal tube layer (irregularly oriented prismatic structure), treated with MUTVEI solution for 5 minutes.

Results

Out of 50 examined Recent and fossil species, tubes of 31 contained a single layer made of mostly isotropic (homogeneous or irregularly oriented, but see below) ultrastructures, while 17 had more than one layer (Table 1). The tube structures of those species that had two or more tube layers involved one layer made of large densely packed crystals and layer(s) composed of relatively sparsely packed fine crystals. In 16 of those 17 species (Table 1) with multilayered tubes, the dense layer was located in the outer tube wall (Figs. 1-2 & 4-5), contrasting with the inner layers made of finer and more sparsely packed crystals (Figs. 3 & 6). The situation is opposite only in the Recent Marifugia cavatica (Figs. 7 & 8), which had less dense layer positioned in the outer tube wall. Tubes of two species (of 31 single-layered species), Placostegus tridentatus and Vitreotubus digeronimoi, were unusual in having the entire tube walls made of large densely packed crystals of relatively same size and arrangement.

The dense outer layers (DOL) made of large crystals were arranged both into oriented (spherulitic prismatic, simple prismatic, regularly ridged prismatic) and unoriented (homogeneous angular crystal, and irregularly oriented platy) structures according to VINN *et alii* (2008b, d) (Figs. 1-2 & 4-5 ; Table 1). DOLs appeared transparent in tubes of Recent species. This transparency is presumably caused by large size of crystals, especially when crystals are densely packed with little organic material between them. However, uniform crystal orientation probably also supports optical transparency. Transparent outer layers varied in thickness, ranging from very thin as found in Crucigera websteri (1/15 of tube wall thickness) to very thick as in Ditrupa (> 1/2 of tube wall thickness). Moreover, tubes of two species, Placostegus tridentatus, and Vitreotubus digeronimoi that have completely transparent tubes were made entirely of densely packed large crystals arranged into simple prismatic structure.

Serpulids with dense tube outer layer (oriented: spherulitic prismatic) are known from the Cretaceous in *Pentaditrupa subtorquata* (Table 1) and from the Eocene in *Pyrgopolon* cf. *mellevillei* (Figs. 1-3) and in *Rotularia spirulae*.

Sparsely packed fine crystals of inner tube layers were organized into isotropic (irregularly oriented prismatic, spherulitic irregularly oriented prismatic and fine homogeneous angular crystal) structures according to VINN *et alii* (2008b) (Fig. 3 & 6).

Discussion

Our study found that different types of serpulid tube ultrastructures are not randomly located in the serpulid tubes made of two or more layers. In all studied species, except for the freshwater cave-dwelling species *Marifugia cavatica*, more dense tube layers are located in the outer part of the tube wall (Table 1). That *Marifugia* shows a reverse organization of layers could be an adaptation to life in oversaturated in calcium carbonate freshwaters of limestone caves. We hypothesize that the location of the tube layers in other studied serpulids (all marine) may indicate an evolutionary adaptation of serpulid polychaetes to counteract predation and carbonate dissolution.



Figure 7: *Marifugia cavatica*, Hercegovina, Popovo Polje, Recent. Polished cross section through outer sparsely packed tube layer, treated with 1% acetic acid for 1 minute.



Figure 8: *Marifugia cavatica*, Hercegovina, Popovo Polje, Recent. Polished cross section through inner dense tube layer, treated with 1% acetic acid for 1 minute.

The calcareous exoskeletons of many marine invertebrates, such as molluscs (TAYLOR & KEN-NEDY, 1969), brachiopods (WILLIAMS & MACKAY, 1978), and bryozoans (BANTA, 1968) are externally covered by an organic layer. This layer isolates mineral parts of the calcareous structures from sea water and could protect against carbonate dissolution and predation (TAYLOR & KENNEDY, 1969; HARPER, 1997; KAR-DON, 1998). Serpulid tubes lack an external organic cover and are in direct contact with seawater that can dissolve structures made of calcium carbonate. The rate of such dissolution in coastal environments, although normally insignificant, is inversely related to temperature (LANGMUIR, 1997). Pressure alone does not affect the calcite solubility as much as temperature does, but when pressure is significant, its effect alone can increase calcite solubility about two fold (KRAUSKOPF & BIRD, 1995). In the deep sea, calcium carbonate accumulation is costly because of low temperatures and tremendous pressures. Dissolution of biogenic carbonate begins at several hundred meters and increases

only gradually until the lysocline, a zone of rapid dissolution. Below the lysocline, carbonate compensation depth (CCD, 4200-5000 m) is reached at which the rate of carbonate accumulation equals the rate of carbonate dissolution, thus, calcareous skeletons of marine animals accumulate at the same rate at which they dissolve (LANGMUIR, 1997; THURMAN & TRUJILLO, 2004). Serpulid inhabit even the deepest parts of the Ocean (KUPRIYANOVA *et alii*, 2011), where carbonate dissolution is intense, and should presumably need special adaptations, both physiological and structural, against carbonate dissolution of their tubes.

Although serpulids are protected by the tubes and the opercula that block the tube entrance, they are attacked by drilling and crushing predators, such as crabs (POLOCZANSKA *et alii*, 2004), echinoids, asteroids, fish (BOSEN-CE, 1979), and especially naticid gastropods (ten HOVE, 1994; TAN & MORTON, 1998; MORTON & HARPER, 2009).



Figure 9 : Phylogenetic tree of serpulids (cladogram of maximum parsimony analysis; modified after KU-PRIYANOVA & NISHI, 2010). Genera studied here indicated with colors. Bold = taxa with two or more tube layers. Red = taxa with outer dense tube layer. Yellow = outer sparsely packed fine crystal tube layer. P = oriented prismatic structures; HA = homogeneous angular crystal structure. *Pomatoceros* in meantime = *Spirobranchus*.

The outer organic layers (conchiolin or periostracum) of bivalves are widely believed to be an adaptation that inhibits drilling by predatory naticid gastropods (see KARDON, 1998 and references therein). Such predation may be similarly discouraged by harder external mineral tube layers of serpulids. The most plesiomorphic serpulid tubes presumably had aragonitic irregularly oriented prismatic (IOP) ultrastructure (VINN et alii, 2008b, c) composed of fine unoriented prismatic crystals in an organic matrix, similar to that known for "Serpula" etalensis from the Early Jurassic (VINN et alii, 2008c). The earliest signs of drilling predation on serpulids (Placostegus) presumably by gastropods are reported from the Late Cretaceous (JÄGER & Koči, 2007; see also ten Hove, 1994). The earliest serpulids with two-layered tubes, containing a dense outer layer and less dense inner layers occur in the Cretaceous (Pentaditrupa subtorquata: VINN, 2005). It appears that in addition to dissolution risk by sea water, serpulids have had to tolerate predation pressure throughout most of their evolutionary history. The appearance of species with DOLs in the Mesozoic could reflect an adaptation against predation because this era, particularly its Jurassic and Cretaceous periods, was a time of increased predation intensities, especially in the shallow water habitats (OJI, 1996). This event, termed by VERMEIJ (1977) as Mesozoic Marine Revolution, may well be associated with the appearance of presumable defensive tube structures in serpulids.

Our study showed that the distribution of Recent serpulid tubes with and without DOLs does not seem to show any ecological pattern, as animals with DOLs are found in all habitats (Table 1). The explanation of the observed pattern appears to be found in the evolutionary history of the group. Phylogenetic studies by KUPRIYANOVA et alii (2006, 2009) and KUPRIYANO-VA & NISHI (2010) consistently indicate that serpulids are split into two major clades A and B (Fig. 9). Serpulid taxa in Clade A are characterized by the most advanced and complex oriented tube ultrastructures such as lamelloordered fibrillar, and fibrillar. spherulitic lamello-fibrillar structures (VINN et alii, 2008b). Clade B is characterized by isotropic (non-oriented) structures, mostly by irregularly oriented prismatic and spherulitic irregularly oriented prismatic structures. Formation of these structures requires less biological control over the mineralization process than formation of complex oriented ultrastructures in the tubes of clade A (VINN et alii, 2008b). Thus, taxa in Clade A have a more advanced biomineralization system than those in the Clade B, the latter appear to be more close to the primitive plesiomorphic condition of serpulid biomineralization system. Although DOLs in the serpulid tubes occur in both major clades of serpulids (Fig. 9), the majority of them (16 out of 17) are found in the apomorphic clade A that contains the most speciose and ecologically successful shallowwater serpulid genera such as Hydroides and Serpula. DOLs are absent in several unusual species of that clade, such as brackish-water species of the genus Ficopomatus, fresh-water Marifugia cavatica, and coral-embedded Floriprotis sabiuraensis and Hydroides spongicola. The external tube walls of three other species from clade A lacking DOLs, Serpula vermicularis, S. israelitica and Spirobranchus triqueter, are composed of complex oriented structures that also have enhanced mechanical properties as compared to the isotropic structures of clade B (VINN et alii, 2008b). Placostegus tridentatus and Vitreotubus digeronimoi, the only taxa with tubes exclusively made of dense structures, also belong to Clade A.

The only taxon in plesiomorphic Clade B having two-layered tubes with DOL is the abyssal Bathyvermilia langerhansi. The dense outer layer in this species is very similar in structure to that found in the tube of Laminatubus alvini, a deep-sea species from clade A. Thus, it is possible that similar (homogeneous angular crystal structure) dense outer layers in tubes of Bathyvermilia and Laminatubus have independently evolved as adaptations to counteract tube dissolution in abyssal depths. In abyssal depth sea water is under saturated of calcium carbonate and skeletons made of CaCO₃ are being dissolved much faster than in shallower depths due to cold temperature and high pressure.

Thus, we suggest that DOLs evolved as a response to increased predation pressure in relatively shallow-water serpulids of the apomorphic clade A, and as a separate adaptation to calcium carbonate dissolution in deep-sea representatives of the plesiomorphic clade B. A detailed phylogeny of the group and additional mineralogical studies involving tubes of abyssal serpulid species are needed to test the hypothesis.

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