



**The lower Aptian ammonites of the Les Ferres Aptian Basin
(Lower Cretaceous, Southeast of France)
Part I: Introduction and biostratigraphy**

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Abstract: On the southern margin of the Vocontian Basin (SE France), for taphonomic reasons (fragmentation, reworking, pyritization), lower Aptian (Lower Cretaceous) deposits of the interval between the *Deshayesites deshayesi* and *Dufrenoyia furcata* zones are usually not suitable for studying the late ontogenetic developments of ammonites (fossil cephalopods). In the vicinity of the village of Les Ferres (Department of the Alpes-Maritimes, SE France), a relatively thick sedimentary succession with well-preserved ammonites, representative of the zones mentioned above, is found in a small basin called the "Les Ferres Aptian Basin" (LFAB). The 1262 ammonites collected or studied *in situ* are dated from the *Deshayesites forbesi* Zone through the top of the *Dufrenoyia furcata* Zone. This study presents the lithostratigraphy, biostratigraphy derived from these ammonites, and the respective range of these taxa. This is an introductory contribution to any future palaeontological study of the lower Aptian ammonites in the Les Ferres area. The lithologic unit overlying the Hauterivian-Aptian limestones is introduced herein as the Les Graous Formation. It is subdividing in three members, from bottom to top: 1) the Combe de Joinet Member, 2) the Pont de la Cerise Member, and 3) the Les Graous Member. In addition, two remarkable levels are identified: the *Ammonitoceras* level (outstandingly abundant) and the *Toxoceratoides* bed.

Key-words:

- Southeast of France;
- lower Aptian;
- biostratigraphy;
- ammonites

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Résumé : *Les ammonites de l'Aptien inférieur du bassin aptien de Les Ferres (Crétacé inférieur, SE de la France). I^{ère} partie : Introduction et biostratigraphie.*- Sur la marge méridionale du Bassin Vocontien (SE de la France), les dépôts d'âge aptien inférieur (Crétacé inférieur) de l'intervalle situé entre les zones à *Deshayesites deshayesi* et *Dufrenoyia furcata* ne sont habituellement pas favorables à l'étude des développements ontogénétiques tardifs des ammonites (céphalopodes fossiles) pour des raisons taphonomiques (fragmentation, remaniement, pyritisation). Dans les environs du village de Les Ferres (département des Alpes-Maritimes, SE de la France), des sédiments aptiens relativement épais et riches en ammonites bien conservées relevant de l'intervalle stratigraphique précité se sont déposés dans un petit bassin structural dénommé ici bassin aptien de Les Ferres (LFAB). Les 1262 ammonites collectées ou étudiées *in situ* sont datées de la Zone à *Deshayesites forbesi* jusqu'au sommet de la Zone à *Dufrenoyia furcata*. La lithostratigraphie, la biostratigraphie fondée sur les ammonites et la distribution de leurs taxons respectifs sont décrits dans ce travail qui constitue une introduction litho- et biostratigraphique à de futurs travaux paléontologiques sur les ammonites de l'Aptien inférieur du secteur des Ferres. Au-dessus des calcaires hauteriviens à aptiens, une nouvelle formation lithologique est décrite, la Formation de Les Graous. Elle est divisée en trois membres, de bas en haut : 1) Membre de la Combe de Joinet, 2) Membre du Pont de la Cerise et 3) Membre de Les Graous. Deux niveaux remarquables y sont identifiés : le niveau à *Ammonitoceras* (exceptionnellement abondants) et le banc à *Toxoceratoides*.

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Mots-clefs :

- sud-est de la France ;
- Aptien inférieur ;
- biostratigraphie ;
- ammonites

1. Introduction

In the Vocontian Basin (Southeast of France), upper lower Aptian ammonite faunas (middle *Deshayesites deshayesi* to *Dufrenoyia furcata* zones sensu BERSAC *et al.*, 2012; see Fig. 1) have been repeatedly described and figured (SAYN, 1920; THOMEL, 1963, 1964, 1968; DAUPHIN, 2002; DUTOUR, 2005; JOLY & DELAMETTE, 2008). Nevertheless, their study is limited by several factors:

- they are generally preserved as small pyrite nuclei, which makes the study of their ontogeny difficult;
- heteromorph ammonites are generally under-represented in the Vocontian Basin for taphonomical reasons (DAUPHIN, 2002; DUTOUR, 2005) and because they were not adapted to a pelagic environment (WESTERMANN, 1996; LUKENEDER, 2015);
- the conditions are usually unfavorable for preservation of large heteromorph ammonites in the neritic margins of the Vocontian Basin, with the upper lower Aptian either being absent or represented by reworked deposits yielding only fragmented ammonites (BERSAC *et al.*, 2010; PICTET *et al.*, 2015).

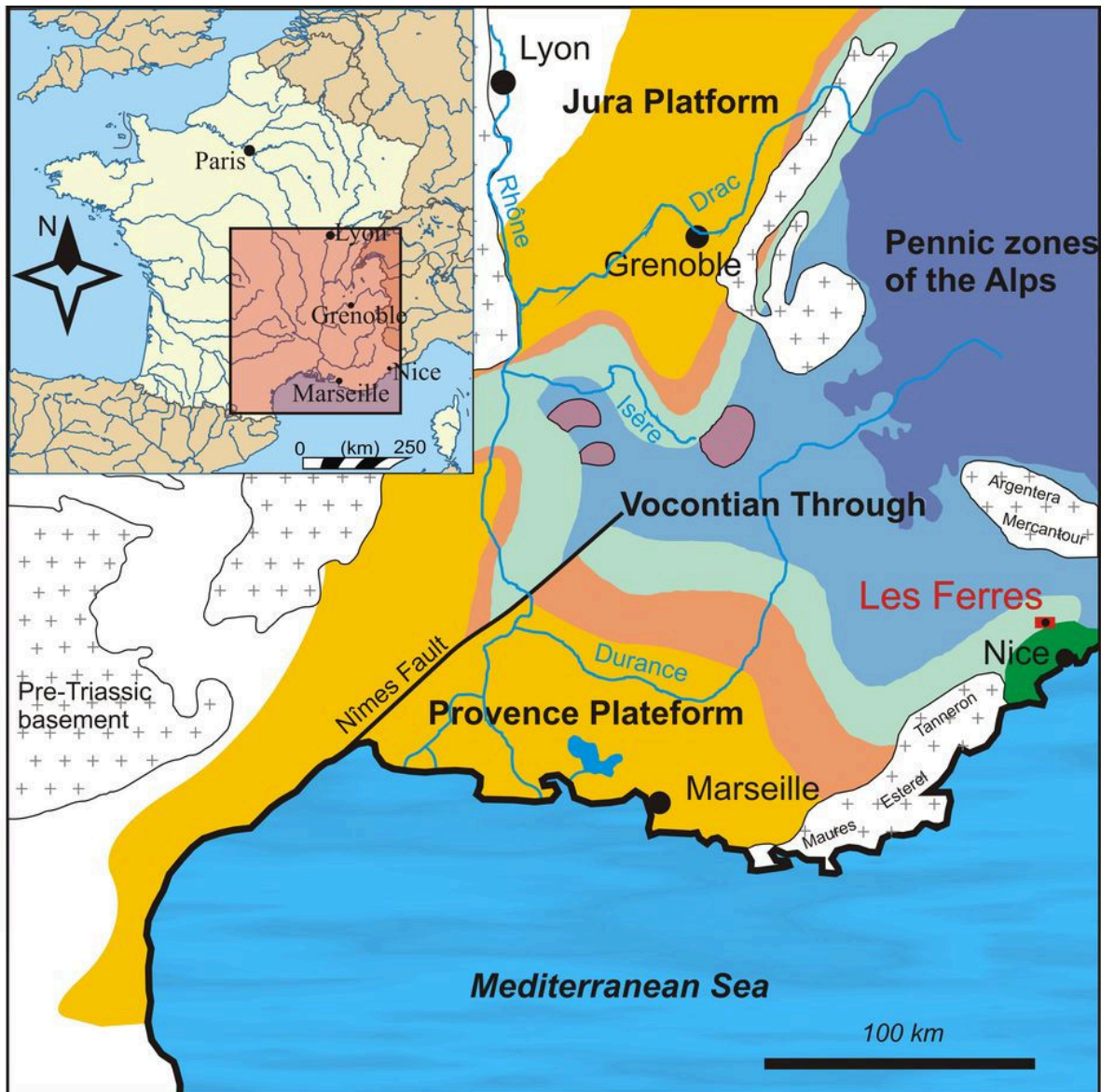
Most of the large lower Aptian heteromorph ammonites in the Vocontian Basin that have been figured in the literature are older than the middle *D. deshayesi* Zone (sensu BERSAC *et al.*, 2012; see Fig. 1 and ROCH, 1927; DELANOY, 1998; DELANOY *et al.*, 2008; PICTET *et al.*, 2009, 2015). A few of the more recent species that have been reported include *Ammonitoceras ucetiae* DUMAS, 1876 (DUMAS, 1876) and *Toxoceratoides?* sp. inc. "Gigantomorph" *godeti* THIEULOY, 1990 [THIEULOY, 1990 = *Lithancylus godeti* (THIEULOY, 1990), see CONTE, 1999 and MIKHAILOVA & BARABOSHKIN, 2001] from the Serviers-et-Labaume and Montaren-et-Saint-Médiers area (Gard department, south-eastern France; DUMAS, 1876). These ammonites come from the *Deshayesites multicostratus* Subzone (*D. deshayesi* Zone sensu BERSAC *et al.*, 2012) based on the evolutionary characters of the associated Douvilleiceratinae PARONA & BONARELLI, 1897 (see BERSAC & BERT, 2018).

The Les Ferres area is located in the lower Esteron Valley (Alpes-Maritimes Department, SE France), in the southern part of the Vocontian Basin. It was situated at the boundary between the neritic and hemipelagic domains during the early Aptian (Figs. 2 - 3). The Aptian series were deposited in a small E-W structural basin, within

a series of tilted blocks. This palaeostructure, which is named herein the Les Ferres Aptian Basin (LFAB), was the site for the deposition of a relatively thick sedimentary sequence with numerous large ammonites. The interval from the

		SZS (Reboulet <i>et al.</i> , 2018)		(Bersac <i>et al.</i> , 2012)	
St.	Zones	Subzones	Zones	Subzones	
Upper Aptian	<i>Epicheloniceras (pars) martini</i>	<i>Epicheloniceras buxtorfi</i>	Not studied	Not studied	Not studied
		<i>Epicheloniceras gracile</i>			
		<i>Epicheloniceras debile</i>			
Lower Aptian	<i>Dufrenoyia furcata</i>	<i>Dufrenoyia dufrenoyi</i>	<i>Dufrenoyia furcata</i>	<i>Dufrenoyia dufrenoyi</i>	
		<i>Dufrenoyia furcata</i>		<i>Dufrenoyia furcata</i>	
	<i>Deshayesites deshayesi</i>	<i>Deshayesites grandis</i>	<i>Deshayesites deshayesi</i>	<i>Deshayesites grandis</i>	
				<i>Deshayesites multicostratus</i>	
				<i>Deshayesites deshayesi</i>	
	<i>Deshayesites forbesi</i>	<i>Roloboceras hambrovi</i>	<i>Deshayesites forbesi</i>		
		?		<i>Deshayesites forbesi</i>	
				<i>Deshayesites fittoni</i>	
	<i>Deshayesites ogilansensis</i>	<i>Deshayesites luppovi</i>	<i>Deshayesites fissicostatus</i>	<i>D. fissicostatus obsoletus</i>	
		?		<i>D. fissicostatus fissicostatus</i>	
			?		

Figure 1: Biostratigraphic charts used in the present work. SZS: Standard Mediterranean Zonal Scheme. St.: stages.









- | | |
|---|---|
|  Deep open marine environment (pelagic) |  Drowned platform (glaucopit condensed series) |
|  Shallow open marine environment (hemipelagic) |  Inner platform |
|  External platform |  Gravity reworked deposits |

Figure 2: Location of the Les Ferres area in the Vocontian Basin (figure after BERSAC *et al.*, 2010, amended).

Deshayesites forbesi to the *Dufrenoyia furcata* zones does not include any polyzonal or condensed intervals. The ammonite fauna is represented by the "classical" taxa described and figured in the pelagic domain (*Dufrenoyia* KILIAN & REBOUL, 1915, *Chelonicerus* HYATT, 1903, *Pseudohaploceras* HYATT, 1900, *Toxoceratoides* SPATH, 1924 and *Colombicerus* SPATH, 1923, etc., see for example DUTOUR, 2005) associated with numerous heteromorph ammo-

nites that belong mainly to the genera *Ammonitoceras* DUMAS, 1876, *Lithancylus* CASEY, 1960, and *Tropaeum* J. de C. SOWERBY, 1837, which were previously poorly known in SE France. This material is sufficiently abundant and stratigraphically constrained to allow the study of evolutionary trends and intraspecific variability. In a nutshell, the LFAB is a unique locality to study ammonites of this time interval.

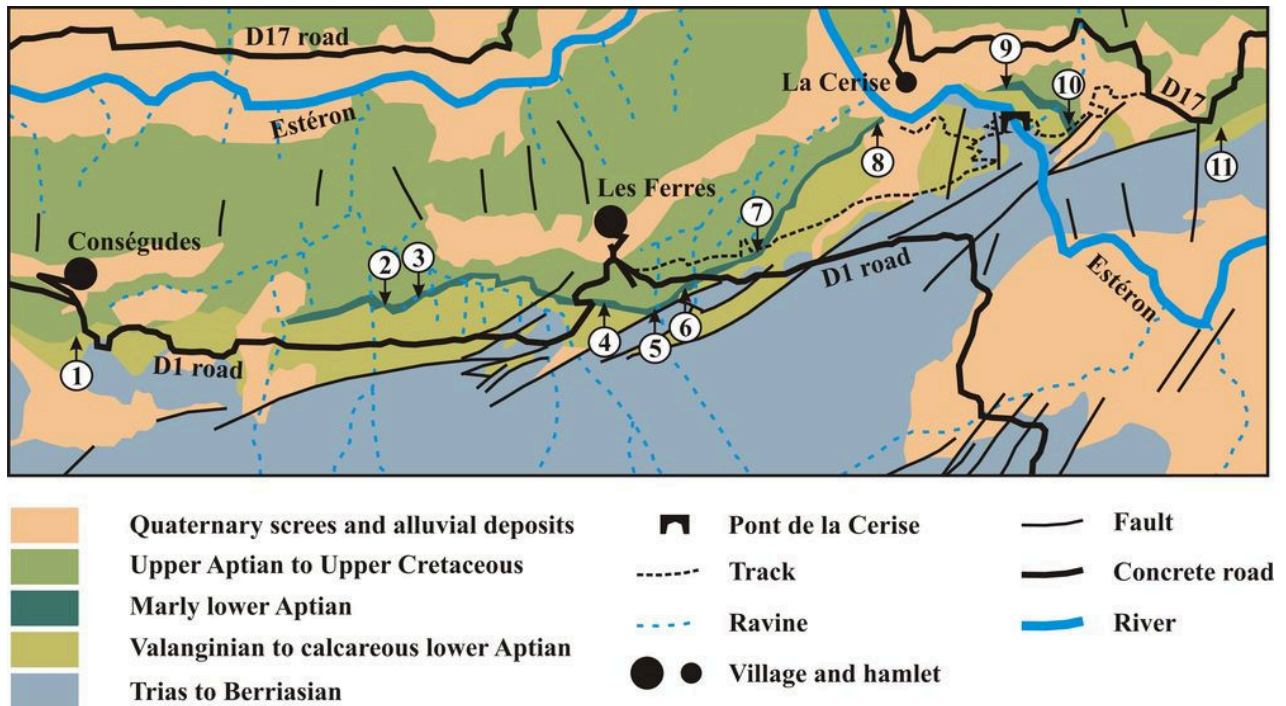


Figure 3: Map of the Les Ferres area with location of the studied or mentioned sections. 1: Conségudes section; 2: Les Graous 2 section (GRS 2); 3: Les Graous 1 section (GRS 1); 4: unnamed section of the La Vallière locality studied by RAGAZZI (1982, p. 73), then BRÉHERET (1997, p. 267, Figs. 114-115); 5: La Vallière section (VAL); 6: La Graou section (GRO); 7: Combe de Joinet section (CHP); 8: Pont de la Cerise section (CLE); 9: Combe de Marin section (CRS); 10: Pont Pairé section (PPR); 11: Colle-Belle section.

The objective of the present work is to describe the lithostratigraphy and biostratigraphy of the lower Aptian of LFAB, with the aim of supporting our future palaeontological studies of ammonites from this area.

2. Palaeogeographic, structural and stratigraphic framework

The studied outcrops all are located on the northern flank of Cheiron Mountain, in the vicinity of the village of Les Ferres (Fig. 3).

The study area belongs to the southern subalpine Castellane Arc, at the southern margin of the Vocontian Basin. It is part of a tectonic structure on the southern boundary of the Estéron Syncline, at the junction between two tilted blocks. This tectonic structure is characterized by the presence of WNW-ESE oriented half-grabens. The (syndimentary) tectonics were active from the Valanginian to the Albian, and are responsible for the rapid lateral change in thickness and the numerous gaps that characterize the Lower Cretaceous succession in this area (DARDEAU, 1987; DARDEAU & GRACIANSKY, 1987). This tectonic structure has also resulted in the formation of a small EW-oriented basin, in which the lower Aptian sediments were deposited. The Lower Cretaceous strata were deposited in the outer neritic domain. Their facies are intermediary between those of the hemipelagic units located to the northeast and those of the condensed neritic series located to

the southeast (COTILLON, 1971; BERSAC *et al.*, 2010; Fig. 2).

The Aptian of the LFAB is represented by two main lithological units: a calcareous unit overlain by a plurimetric marly unit (COTILLON, 1971, 2010; RAGAZZI, 1982; Fig. 4); the calcareous unit and the lower part of the marly unit are referred to the lower Aptian and the rest of the marly unit to the upper Aptian. These deposits form a transgressive sequence (COTILLON, 1971) related to the "*middle Bedoulian drowning event*" (BRÉHERET, 1997; COTILLON, 2010; MASSE & FENERCI-MASSE, 2011, p. 671). The present work focuses on the marly lower Aptian of the LFAB. Newly identified lithostratigraphic units are formally described in Chapter 5.2.

3. Historical account of the Aptian ammonite palaeontology in the Les Ferres area

FALLOT (1885, p. 128) was the first author to indicate the possible presence of Aptian strata in the vicinity of Les Ferres. Subsequently, GOGUEL (1944), COTILLON (1971), GINSBURG *et al.* (1980), RAGAZZI (1982), and BRÉHERET (1997) have documented the lithostratigraphy and ammonite biostratigraphy of this area (see detailed historical accounts in RAGAZZI, 1982, and BRÉHERET, 1997). COTILLON (1971, p. 146), for instance, was the first to mention the presence of the ammonite genus *Dufrenoyia* (and consequently the presence of the *D. furcata* Zone, see Fig. 1) in the LFAB. However,



several occurrences of ammonite taxa reported in these contributions were not supported by any illustrations. Our own investigations in the area started in the early 2000s and resulted in preliminary publications (BERT & BERSAC, 2011; BERT, 2014) including illustrations of few specimens. More recently, DELANOY *et al.* (2018) revisited the site. They figured and described these faunas in detail. These latter authors, focused on large heteromorphic ammonites and described ten new taxa.

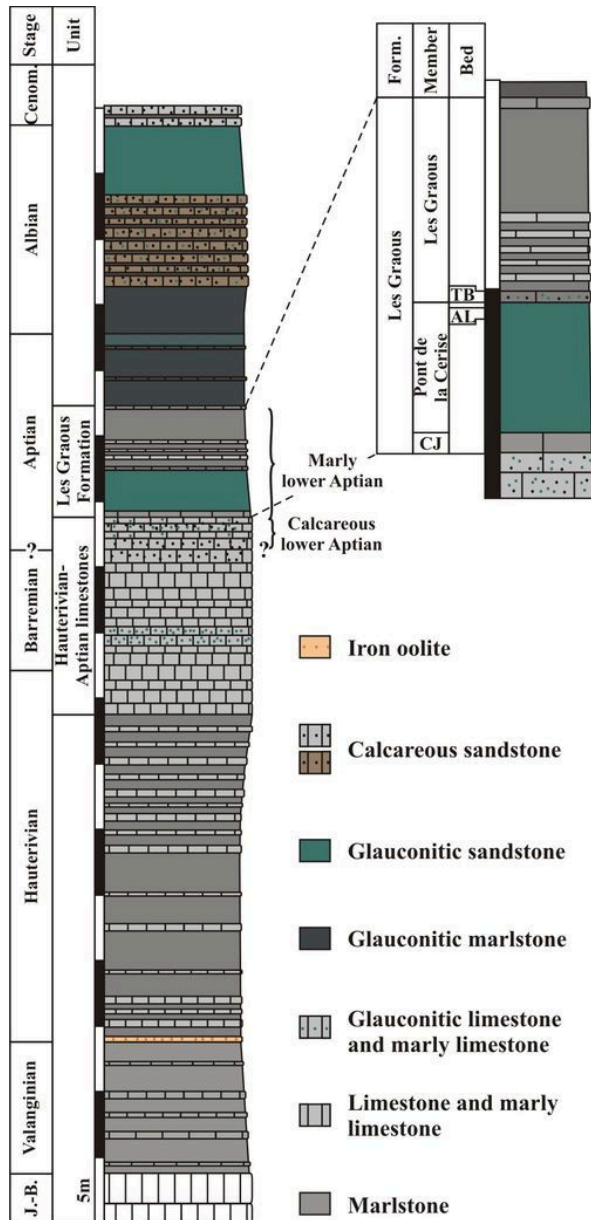


Figure 4: Synthetic log of the Lower Cretaceous of the Les Ferres area. J.-B.: Jurassic and Berriasian; Cenom.: Cenomanian; Form.: Formation; CJ: Combe de Joinet Member; AL: *Ammonitoceras* Level; TB: *Toxoceratoides* Bed.

4. Methods

4.1. Lithostratigraphic units and sections

To date, the lithostratigraphic units of the LFAB have not been formally described in the scientific literature. In the present work, we propose a formal lithostratigraphic scheme for the marly lower Aptian of the LFAB (Fig. 4 and see Chapter 5.2). The name of the recognized units is that of the outcrop where they are visible under the best conditions. Two remarkable levels are recognized and named according to the characteristics of their respective ammonite faunas.

4.2. Biostratigraphic scheme

The biostratigraphic scheme used in the present work is that previously proposed by BERSAC *et al.* (2012) for the lower Aptian, based on a review of the Deshayesitidae STOYANOW, 1949 (Ammonoidea, BERSAC & BERT, 2012, 2015; Fig. 1).

4.3. Field sampling and material studied

Our objective was to prospect all the Aptian outcrops of the LFAB. Nine sections were then identified and 8 of them were sampled (Fig. 3). Two other sections close to the study area were also studied: the Conségudes and the Colle-Belle sections. A total of 1,262 ammonites was bed-by-bed sampled or studied *in situ* between 2000 and 2017. This material is deposited in the collections of the Laboratory of the Research Group for Paleobiology and Biostratigraphy of the Ammonites (G.P.A., France).

4.4. Remarks about the material published by DELANOY *et al.* (2018)

We were both present in the field when all the material published by DELANOY *et al.* (2018) was extracted and we reviewed it once prepared. Therefore, we deemed it necessary to include it in our study. Our own field data differ somewhat from those of DELANOY *et al.* (2018), and we will explain these differences below where necessary. It should be noted that the designation of specimens published by DELANOY *et al.* (2018) is complicated by the fact that several specimens may bear the same number, or that several numbers may be assigned to the same specimen. The problematic numbers are 28754, 28756, 28767, 28766, 28768, and 28769. Similarly, it should be noted that the holotype of *Proaustraliceras bournaudi* DELANOY *et al.* 2018, numbered AP-001 (DELANOY *et al.*, 2018, p. 47 and Fig. 7) does not belong to the BERSAC's collection contrary to the statement by these authors.



5. Results

5.1. PRELIMINARY AMMONITE TAXONOMIC CONSIDERATIONS

In this chapter, we briefly discuss the taxonomic assignments used in our study for the sole purpose of providing support for the presentation of the lithostratigraphic and biostratigraphic results. In particular, according to our data, the new taxa introduced by DELANOY *et al.* (2018) are all subjective synonyms of taxa already described. For this reason, these taxa are not quoted in the stratigraphic logs (Figs. 5, 7, 9 - 10, 12, 14 - 15). We will expose our taxonomic positions in detail in future works.

5.1.1. The Deshayesitidae and Douvilleiceratidae PARONA & BONARELLI, 1897

These two ammonite families provide lower Aptian zonal and subzonal index species for different palaeogeographic domains (CASEY *et al.*, 1998; BARABOSHKIN & MIKHAILOVA, 2002; REBOULET *et al.*, 2018). The criteria for identifying Deshayesitidae used in this work are those proposed by BERSAC & BERT (2012, 2015). They rely on the morphology of the ventral area and the suture lines. The taxa we identified in the LFAB are *Deshayesites* sp. (Pl. 1, figs. 4-5), *Deshayesites multicosstatus* SWINNERTON, 1935 (Pl. 2, figs. 6-8, 13-15, 19), *Deshayesites grandis* SPATH, 1930 (Pl. 3, figs. 9-12), *Dufrenoyia furcata* (J. de C. SOWERBY, 1837) (Pl. 5, figs. 5-9), and *Dufrenoyia dufrenoyi* (ORBIGNY, 1841) (Pl. 5, figs. 30-32). The lower Aptian Douvilleiceratidae of the LFAB are only represented by Douvilleiceratinae. Their identification is based on the duration of their ontogenetic stages (see BERSAC & BERT, 2018, and references therein). The identified taxa are *Procheloniceras* sp. (Pl. 1, figs. 6-7), *Cheloniceras* sp. and *Cheloniceras crassum* SPATH, 1930 (Pl. 2, figs. 16-17; Pl. 3, figs. 5-6; Pl. 4, figs. 8-9).

5.1.2. The Parahoplitidae SPATH, 1922

The collected material is assigned to the *D. furcata* Subzone (*D. furcata* Zone) (Pl. 5, figs. 10-12). All intermediates can be observed from slender forms having a *Colombiceras crasscostatum*-type shape to robust forms with a *Gargasicerias gargasensis*-type shape. Consequently we consider *Gargasicerias gargasensis* (ORBIGNY, 1841) as a subjective synonym of *Colombiceras crasscostatum* (ORBIGNY, 1841), as already suggested by JACOB (1907), DUTOIR (2005) and BULOT (2010). Possible primitive representatives of Parahoplitidae occur in the *Deshayesites grandis* Subzone (*D. deshayesi* Zone) and are identified as ?*Colombiceras* sp.

5.1.3. The Ancyloceratidae MEEK, 1876

5.1.3.1. *Ammonitoceras*

The genus *Ammonitoceras* is represented by large ammonites with criocone, aspinoceratic or

ancyloceratic coiling, characterized by the presence of a particular ontogenetic stage in the inner whorls: the *Ammonitoceras* stage. This stage is characterized by the presence of bituberculated main ribs bearing an umbilical tubercle and a lateral tubercle of larger size. In a future article, we will publish a detailed study of *Ammonitoceras* from the *D. deshayesi* and *D. furcata* zones of the LFAB.

Preliminary results from this study suggest that:

1. the genus *Ammonitoceras* shows significant intraspecific variability that mainly relates to the duration of the ontogenetic *Ammonitoceras* stage;
2. *Ammonitoceras* is probably sexually dimorphic, with large crioconic macroconchs (Pl. 2, figs. 1-2; Pl. 4, figs. 1-3) and microconchs with aspinoceratic or ancyloceratic coiling (Pl. 2, figs. 3-5);
3. *Ammonitoceras* macroconchs of the *D. furcata* Zone (Pl. 4, figs. 1-3) have on average an *Ammonitoceras* stage disappearing later during ontogeny and a narrower spiral hiatus than those of the *D. multicosstatus* Subzone (Pl. 2, figs. 1-2);
4. *Ammonitoceras* of the *D. multicosstatus* Subzone can be assigned to the species *Ammonitoceras uctiae* and those of the *D. furcata* Zone to *Ammonitoceras lahuseni* (SINZOW, 1906);
5. we interpret *Ammonitoceras dumasi* DELANOY *et al.*, 2018 as the macroconch of *A. uctiae* and regard it as a subjective synonym of that species;
6. we interpret the following species and specimens as microconchs of *A. uctiae*: *Caseyites esteronsensis* DELANOY *et al.*, 2018, *Caseyites vermeuleni* DELANOY *et al.*, 2018, *Ammonitoceras madouxi* DELANOY *et al.*, 2018, *Proaustraliceras bournaudi* DELANOY *et al.*, 2018, the specimen identified under ?*Proaustraliceras* sp. numbered 28766 or 28767 in DELANOY *et al.* (2018, p. 50-51, Fig. 9.A) and the specimen of *Ammonitoceras* aff. *madouxi* numbered 28764 in DELANOY *et al.* (2018, p. 71-72, Fig. 19);
7. according to our interpretation, *A. leiferransensis* DELANOY *et al.*, 2018, *Caspianites ragazziae* DELANOY *et al.*, 2018, and the specimen figured under ?*Caspianites* sp. and numbered 28744 or 28754 in DELANOY *et al.*, (2018, p. 83, Fig. 25.A-B and p. 84) are all macroconchs of *A. lahuseni* and thus subjective synonyms of that species;
8. in our view, *Caseyites morenobedmari* DELANOY *et al.*, 2018 is a microconch of *A. lahuseni* and consequently a subjective synonym of that species; and
9. *Ammonitoceras* specimens from the *D. grandis* Subzone are too incomplete to be identified at the specific rank.



5.1.3.2. *Lithancylus*

Lithancylus of the *D. multicosatus* Subzone of the LFAB (Pl. 2, figs. 11-12) appear to have significant intraspecific variability, which concerns both their ontogenetic sequence and adult size. In our opinion, these specimens belong to the species *Lithancylus grandis* (J. de C. SOWERBY, 1828), and consequently we interpret *Lithancylus bifurcatus* DELANOY *et al.*, 2018 as a subjective synonym of that species.

5.1.4. The Helicancyliidae HYATT, 1894

On the basis of the diagnostic characteristics proposed by BULOT *et al.* (2017) and FRAU *et al.* (2017), we identified *Toxoceratoides* sp., *Toxoceratoides* aff. *royerianus* (ORBIGNY, 1842) (Pl. 3, figs. 13-16) and *Toxoceratoides rochi* CASEY, 1961 (Pl. 5, figs. 1-4).

5.1.5. The Ptychoceratidae GILL, 1871

The genus *Ptychoceras* ORBIGNY, 1842, is represented by rare specimens we identified as *Ptychoceras emericianum* ORBIGNY, 1842 (Pl. 5, fig. 15) on the basis of the revision work of VERMEULEN *et al.* (2015).

5.1.6. The Desmoceratidae ZITTEL, 1895

We identified the Desmoceratidae of the LFAB on the basis of the diagnostic criteria proposed by DUTOUR (2005), in particular the aspect of the whorl section. In this respect, identification of the collected specimens at the specific rank was made difficult by the fact that they are almost always deformed and crushed. The taxa identified are: *Pseudohaploceras* sp. (Pl. 1, fig. 8), *Pseudohaploceras liptoviensis* (ZEUSCHNER, 1856) (Pl. 2, figs. 9-10; Pl. 4, figs. 6-7), *Pseudosaynella* sp., *Aconeceras nisum* (ORBIGNY, 1841) (Pl. 5, figs. 13-14) and ?*Aconeceras* sp.

5.1.7. The Lytoceratidae NEUMAYR, 1875

Representatives of this family are rare in the LFAB and are attributed to *Lytoceras* sp. (Pl. 4, figs. 10-11).

5.1.8. The Macroscaphitidae HYATT, 1900

Based on BUSNARDO's remarks (in GAUTHIER, 2006), we identified *Macroscaphites* aff. *yvani* (PUZOS, 1832) and *Macroscaphites striatisulcatus* (ORBIGNY, 1841) (Pl. 2, fig. 20; Pl. 3, figs. 3-4; Pl. 5, fig. 16).

Following the works of WRIGHT *et al.* (1996), KLEIN *et al.* (2007) and VAŠIČEK *et al.* (2014), we consider *Pseudocrioceratites* EGOIAN, 1969, as a subjective synonym of *Acantholytoceras* SPATH, 1923. The *Acantholytoceras* we collected in the LFAB are too fragmented to be identified to specific level (Pl. 4, figs. 4-5).

5.1.9. The Phylloceratidae ZITTEL, 1884

In the LFAB, this family is represented only by the subgenus *Hypophylloceras* SALFELD, 1924 (*Phylloceras* SUESS, 1865). JOLY & DELAMETTE (2008, Fig. 74) recognized 6 species of *Phylloceras* (*Hypophylloceras*) ranging through the *D. deshayesi* and *D. furcata* zones (sensu REBOULET *et al.* 2018) in the Vocontian Basin. The diagnostic characters are subtle variations in ornamentation, section or morphology of the suture line. Our specimens are not sufficiently well preserved to be identified to the specific level (Pl. 3, fig. 1; Pl. 4, fig. 12).

5.2. LITHOSTRATIGRAPHIC FRAMEWORK

5.2.1. Top of the Hauterivian-Aptian limestones (Fig. 4)

In the neritic domain of the Castellane Arc, Hauterivian to lower Aptian deposits are represented by a pluridecametric succession of massive limestone beds. COTILLON (1971) divided it into a series of formations (formations 17 to 25, COTILLON, 1971, p. 34 and Fig. 12bis). In the LFAB, the Hauterivian-Aptian limestones are represented by a plurimetric succession of massive decimetric limestone beds separated by centimetric marly intervals (Fig. 4) that have already been described in the literature under various names ("*Calcaire blanc compacte*" in FALLOT, 1885, p. 127, "*Barrémo-Bédoulien*" in RAGAZZI, 1982, p. 73).

The highest beds in the upper part of this lithological unit contain increasing amount of sandstone and glauconite (RAGAZZI, 1982), and corresponds to COTILLON's (1971) Formation 25, and these are assigned to the lower Aptian (*pars*, GINSBURG *et al.*, 1980; RAGAZZI, 1982). The upper limit of the Hauterivian-Aptian limestones is represented by a firm-ground-type, with iron crust layers and extensive bioturbation by *Rhizocorallium* ZENKER, 1836 (Fig. 8.D). This discontinuity surface is correlated with the "*early Aptian unconformity*" (COTILLON, 2010, p. 4) that is related to the "*middle Bedoulian drowning event*" (MASSE & FENERCI-MASSE, 2011, p. 671). It corresponds to a large-scale discontinuity surface formed during the flooding of the Provençal platform (BRÉHERET, 1997; COTILLON, 2010; MASSE & FENERCI-MASSE, 2011). This event is assigned to the upper *D. forbesi* Zone (PICTET *et al.*, 2015; Fig. 1).

We only sampled the last Hauterivian-Aptian limestone bed: it is represented by a centimetric level of grey to yellowish sandstone relatively rich in internal moulds of ammonites, nautiloids and belemnites (Fig. 13.A). RAGAZZI (1982, p. 74-75) reported the following ammonite fauna: *Ancyloceras audouli* (ASTIER, 1851), *Deshayesites latilobatus* (SINZOW, 1910) [= '*D. latelobata* (SINZOW)' in RAGAZZI, 1982, p. 75], *Deshayesites* sp., *Chelonicer*

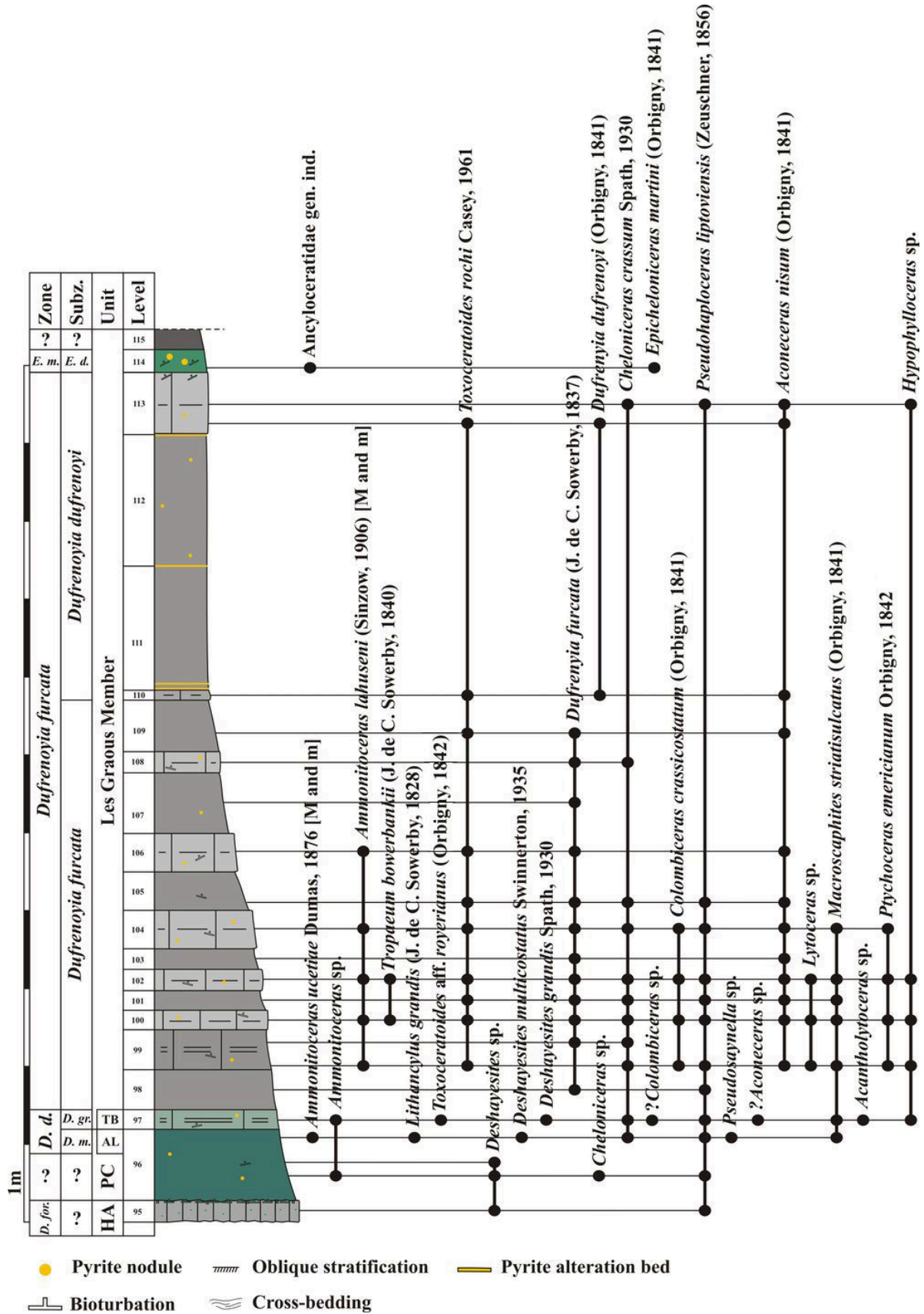


Figure 5: Log of the Les Graous 1 section (GRS 1) with distribution of the ammonite taxa, see Fig. 4 for caption. HA: Hauterivian-Aptian limestones. PC: Pont de la Cerise Member. *D. for.*: *Deshayesites forbesi*. *D. d.*: *Deshayesites deshaysi*. *D. m.*: *Deshayesites multicostatus*. *D. gr.*: *Deshayesites grandis*. *E. m.*: *Epicheloniceras martini*. *E. d.*: *Epicheloniceras debile*. [M]: macroconch, [m]: microconch.

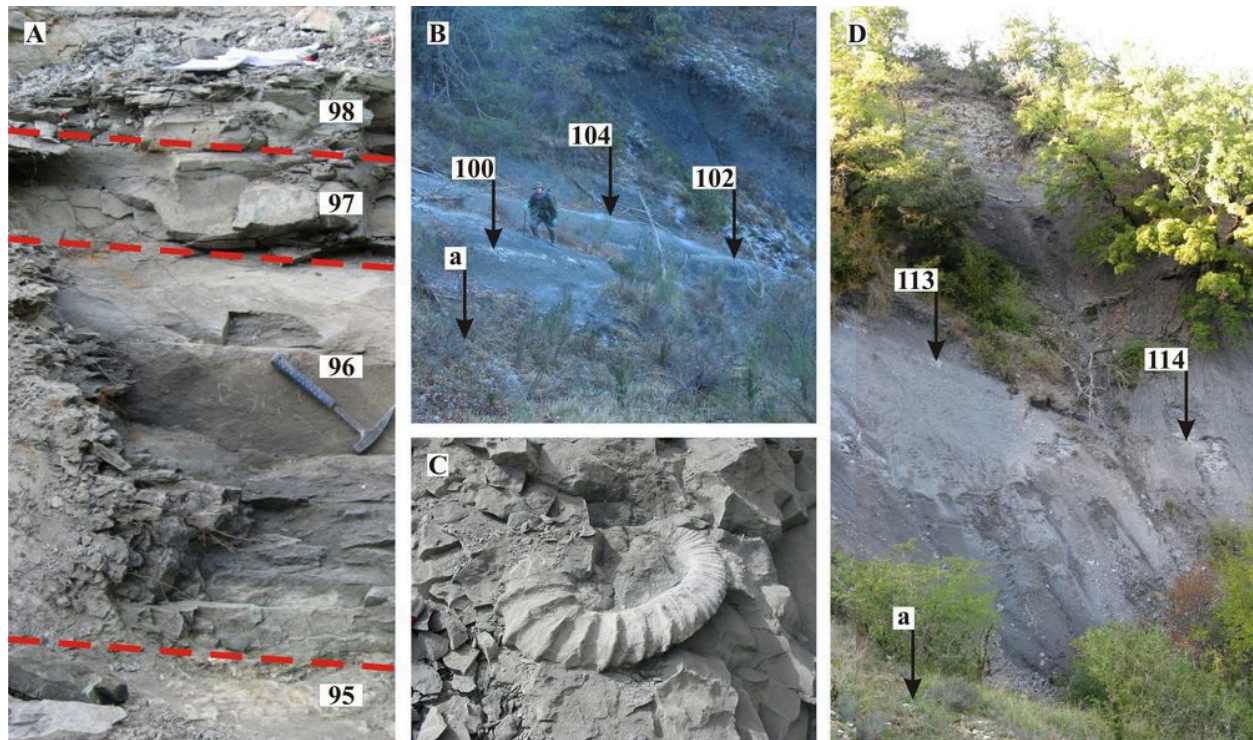


Figure 6: Field photographs of the Les Graous 1 section (GRS 1). Log of this section is figured in Fig. 5. **A:** Beds 95 to 98 at the base of the section in 2009. Length of the hammer: 330 mm. **B:** Beds 100 to 104 in the lower part of the Les Graous Member in 2005. **C:** specimen No. 28750, holotype of *Caseyites morenobedmari* DELANOY *et al.*, 2018 (DELANOY *et al.*, 2018, Fig. 15, = *Ammonitoceras lahuseni* (SINZOW, 1906) [m] in Fig. 5) during extraction from Bed 102 of GRS 1 in 2006. Length of the ammonite: 468 mm (DELANOY *et al.*, 2018, p. 63). **D:** lower Aptian deposits of GRS 1, overlain by upper Aptian and Albian in 2010; a: top of the Hauterivian-Aptian limestones.

sp., *Procheloniceras sp.*, *Pseudohaploceras sp.*, *Costidiscus recticostatus* (ORBIGNY, 1841). We identified: *Ancyloceras rochi* SARKAR, 1955 (Pl. 1, figs. 1-2), *Ancyloceras sp.* (Pl. 1, fig. 3), *Toxoceratoides sp.*, *Deshayesites sp.* (Pl. 1, figs. 4-5), *Procheloniceras sp.* (Pl. 1, figs. 6-7), *Pseudohaploceras sp.* (Pl. 1, fig. 8), *Pseudohaploceras liptoviensis*, *Hypophylloceras sp.* This fauna is associated with numerous nautiloids [*Cymatoceras neckerianus* (PICTET, 1847) (Pl. 1, fig. 9)] and belemnites (*Mesohibolites sp.*).

We could not identify any *Deshayesites* from this bed at the specific rank because the characteristics of their ventral area in the inner whorls and their suture line could not be observed (BERSAC & BERT, 2012, 2015). Nevertheless, their rounded ventral areas on the phragmone suggests that they are older than the *D. grandis* Subzone (Pl. 1, fig. 5).

The Douvilleiceratidae collected in this bed have a relatively long and well expressed *Procheloniceras* stage and no *Cheloniceras* stage was observed (BERSAC & BERT, 2018), but the innermost whorls of the studied specimens are not preserved: if the *Cheloniceras* stage occurs, it is restricted to innermost whorls only (Pl. 1, figs. 6-7). The shape of these specimens with more primitive features than *Cheloniceras cornuelianum* (ORBIGNY, 1841) leads us to attribute them to the genus *Procheloniceras* SPATH, 1923. This genus seems to occur at the top of the *D.*

forbesi Zone (sensu BERSAC *et al.*, 2012), whereas the genus *Cheloniceras* succeeds it from the base of the *D. deshayesi* Zone (BERSAC & BERT, 2018). It should be noted that we have not collected any representative of the stratigraphically significant family Roloboceratinae CASEY, 1961. Their representatives occur in the lower part of the *D. forbesi* Zone (sensu BERSAC *et al.*, 2012) in the Boreal realm and at least from the upper part of the *D. forbesi* Zone or from the base of the *D. deshayesi* Zone (sensu BERSAC *et al.*, 2012) in the Tethysian realm (CASEY *et al.*, 1998; BERSAC & BERT, 2012, Fig. 17, 2015, Fig. 15).

Small *Ancyloceras* ORBIGNY, 1842, similar to those collected at the top of the Hauterivian-Aptian limestones, were reported in the *Deshayesites weissi* Zone (= *D. forbesi* Zone sensu REBOULET *et al.*, 2018) by ROPOLO *et al.* (1998, p. 171) in the Roquefort-la Bédoule area (SE France); they were identified as *Ancyloceras matheronianum* ORBIGNY, 1842. Pierre ROPOLO provided photographs of these specimens to us and, from our point of view, they are similar to the type specimen of *Ancyloceras rochi*, originally described and figured by ROCH (1927, p. 29, Pl. IV, fig. 2) as "*Ancyloceras Binelli ASTIER*".

Based on these considerations, the last Hauterivian-Aptian limestone bed of the LFAB probably belongs to the lower part of the *D. forbesi* Zone sensu BERSAC *et al.* (2012; Fig. 1).



5.2.2. The Les Graous Formation

This formation is formally defined for the first time in the present work (Fig. 4) and is restricted to the LFAB. It is represented by dark green glauconitic sandstones in its lower part and grey-blue marls in its upper part. Its lower limit is the terminal discontinuity of the Hauterivian-Aptian limestones. Its upper limit is located at the top of a marly-limestone level just below a dark grey glauconitic level (Fig. 4). This glauconitic level yielded the oldest *Epicheloniceras martini* (ORBIGNY, 1842) of the study area and was therefore assigned to the upper Aptian *Epicheloniceras debile* Subzone (*Epicheloniceras martini* Zone, see DUTOUR, 2005 and Fig. 1). The Les Graous Formation type section is Les Graous 1 (GRS1, Figs. 3, 5 - 6, and see Chapter 5.3.1), where it is observable in its entirety. The Les Graous Formation is the equivalent of COTILLON's (1971, p. 69) formations ab1 and ab2 of the neritic domain of the Castellane Arc. At Les Ferres, this author (*ibid.*, p. 146) named it "[a, b]".

We propose to divide the Les Graous Formation into several lithostratigraphic members, described here for the first time:

5.2.2.1. The Combe de Joinet Member

This lithostratigraphic unit is restricted to the LFAB (Fig. 4). It was previously mentioned by RAGAZZI (1982, p. 76), who named it "*marnes bleutées*" ("bluish marls").

The Combe de Joinet Member is represented by 1.10-1.50 m of light grey marly limestone in which are sometimes intercalated two more indurated decimetric levels or a glauconitic decimetric level. The lower limit of this member is the discontinuity surface at the top of the Hauterivian-Aptian limestones. Its upper limit is an intensely bioturbated discontinuity surface (ichnogenus *Thalassinoides* EHRENBERG, 1944; Fig. 13.B). The type section of this member is the Combe de Joinet section (CHP, Figs. 3, 10 - 11, and see Chapter 5.3.4). The ammonite fauna of the Combe de Joinet Member is particularly rare and only represented by fragmented internal moulds of *Ancyloceras* sp., *Deshayesites* sp., *Cheloniceras crassum*, *Macroscaphites* aff. *yvani*, *Pseudohaploceras liptoviensis* and *Pseudohaploceras* sp. This member crops out in the following sections: La Vallière (VAL, see Figs. 3, 7 - 8, and Chapter 5.3.2), Combe de Joinet (CHP), Pont de la Cerise (CLE, see Figs. 3, 12 - 13, and Chapter 5.3.5) and Combe de Marin (CRS, see Figs. 3, 14, and Chapter 5.3.6). It also partly crops out in the Pont Pairé section (PPR, see Figs. 3, 15 - 16, and Chapter 5.3.7).

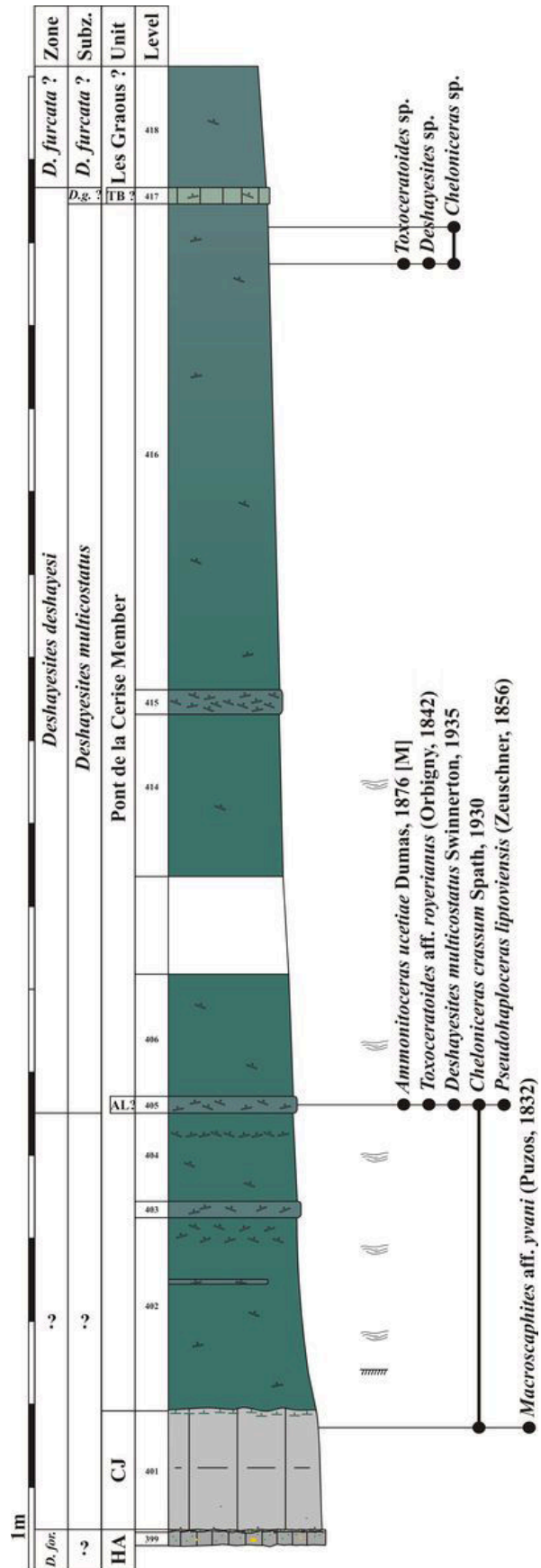


Figure 7: Log of the La Vallière section (VAL) with distribution of ammonite taxa, see Figs. 4 - 5 for caption.

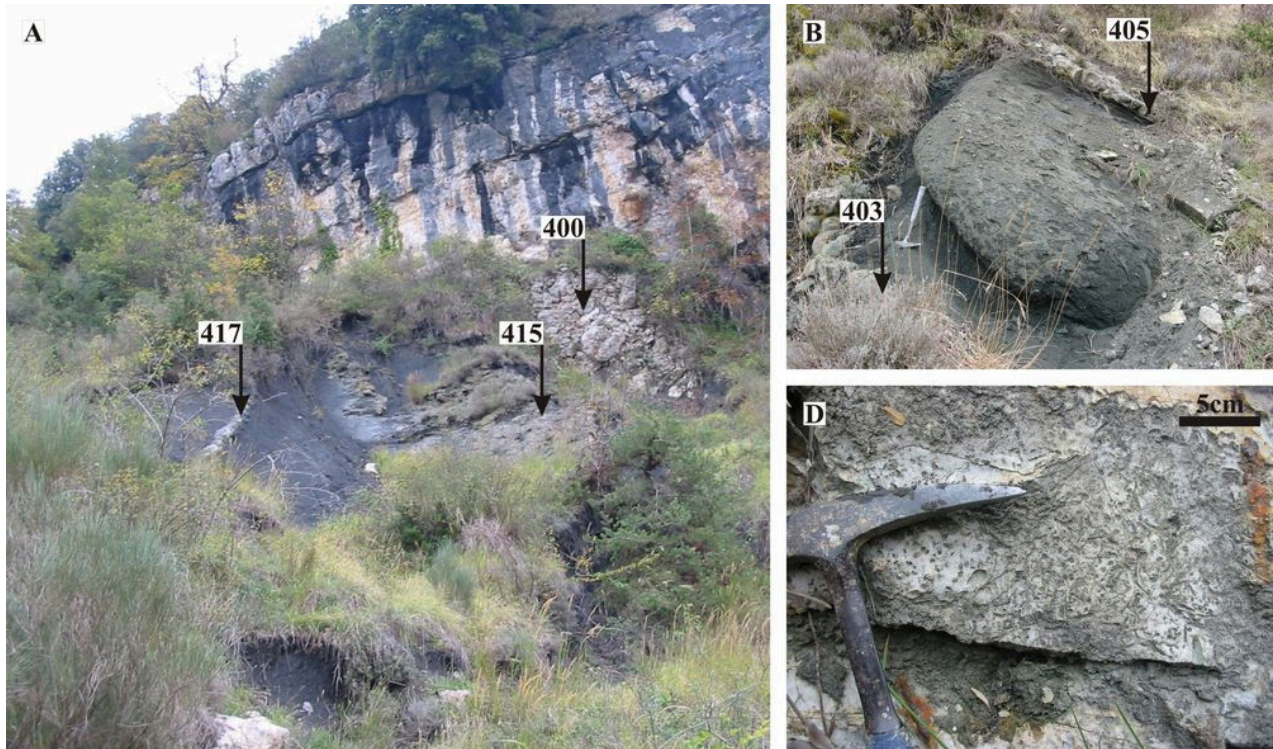


Figure 8: Field photographs of the La Vallière section (VAL). Log of this section is figured in Fig. 7. **A:** general view (2004) of the upper part of the section, located against a palaeo-escarpment. **B:** Beds 403 to 405 of the Pont de la Cerise Member in 2010. Bed 405 is interpreted to be the local equivalent of the *Ammonitoceras* Level. **C:** upper surface of the last bed (400) of the Hauterivian-Aptian limestones with *Rhizocorallium*.

5.2.2.2. The Pont de la Cerise Member

This member is restricted to the LFAB (Fig. 4) and is represented by massive glauconitic sandstones (RAGAZZI, 1982; BRÉHERET, 1997) in which are often intercalated more or less continuous light grey decimetric beds (Figs. 8.A-B, 16.A). The thickness of this member varies from less than one metre in the Les Graous 1 section (GRS1) to 15 m in the La Vallière section (VAL, Fig. 17). The lower limit of this member is represented by the discontinuity surface at the top of the underlying Combe de Joinet Member. Its upper limit is characterized by the abrupt end of the glauconitic sandstones, which are overlain by a characteristic, more or less marly, grey limestone bed, here named the "*Toxoceratoides* Bed" (which belongs to the overlying Les Graous Member, see Chapter 5.2.2.3). The type section of the Pont de la Cerise Member is at the Pont de la Cerise section (CLE, see Figs. 3, 12 - 13, and Chapter 5.3.5). The first decimetres of this member are relatively rich in phosphate nodules, belemnite fragments and fish teeth. When this member reaches a plurimetric thickness, the first two meters display cross-stratifications. Bioturbation is frequent throughout the Pont de la Cerise Member, especially in the more indurated levels (Figs. 5 - 16, *Thalassinoides* in Fig. 13.C).

The glauconitic sandstones of the Pont de la Cerise Member were first mentioned by GOGUEL (1944, p. 21) who assigned them to the Albian. GINSBURG *et al.* (1980) and RAGAZZI (1982, p. 74) dated them as Aptian. A level remarkably rich in ammonites generally occurs in the upper 0.5 m of this member. It is named here the "*Ammonitoceras* Level" and will be described below. The Pont de la Cerise Member crops out in the following sections: Les Graous 1 (GRS1), Les Graous 2 (GRS2), La Vallière (VAL), Combe de Joinet (CHP), Pont de la Cerise (CLE), Combe de Marin (CRS), and Pont Pairé (PPR).

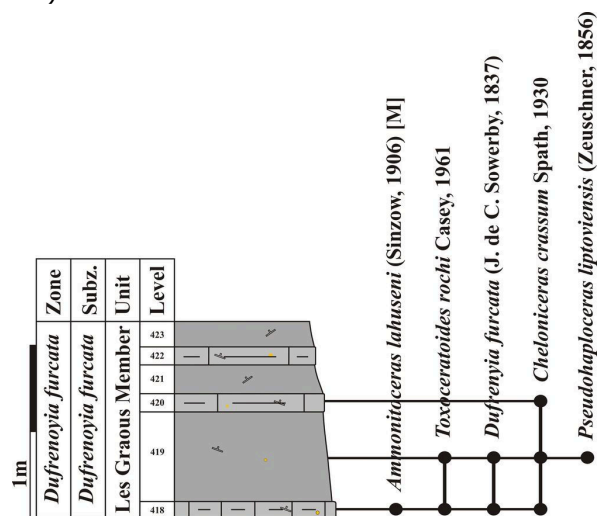


Figure 9: Log of the La Graou section (GRO) with distribution of ammonite taxa, see Fig. 5 for caption.

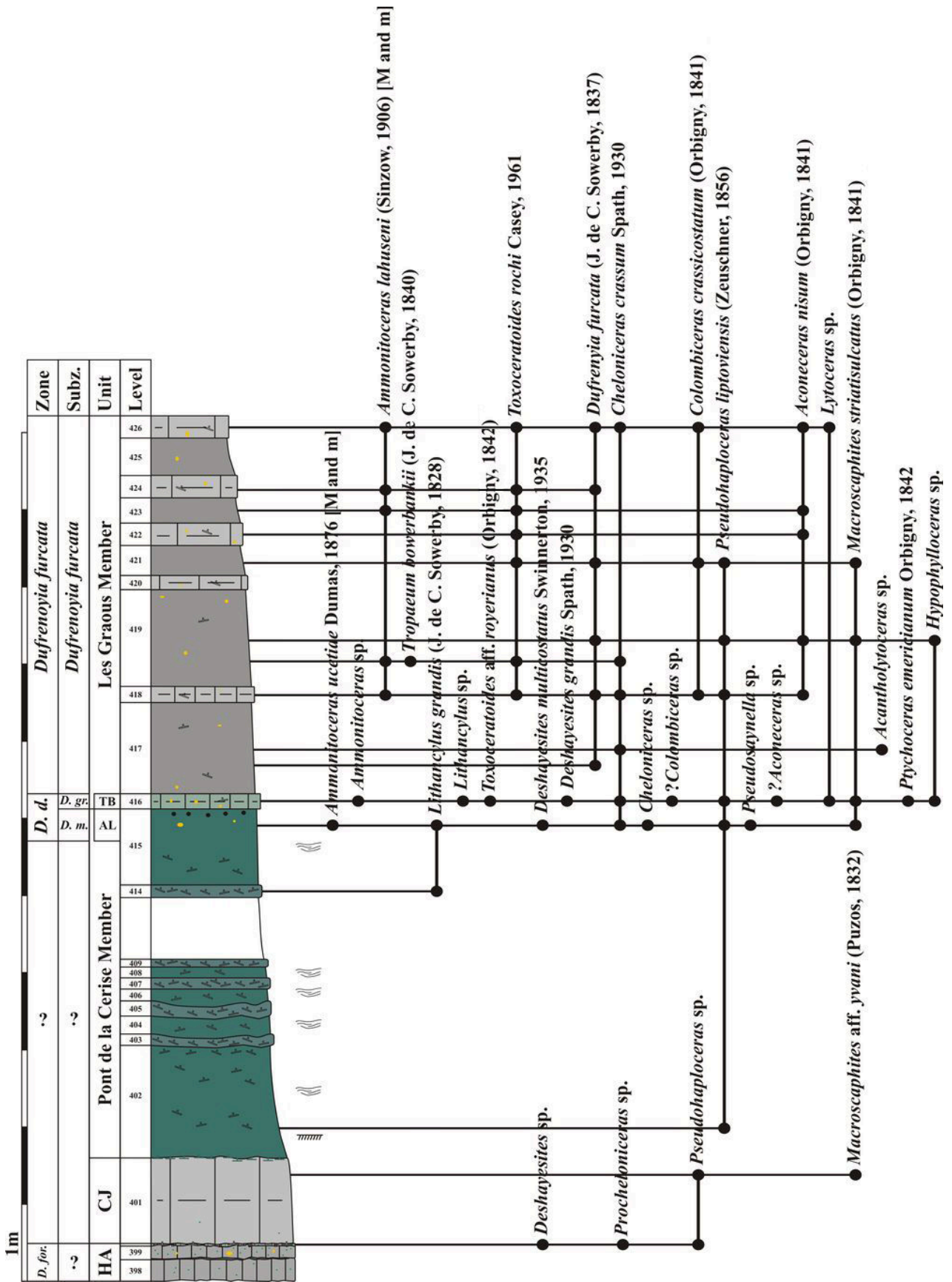


Figure 10: Log of the Combe de Joinet section (CHP) with distribution of ammonite taxa, see Figs. 4 - 5 for caption.

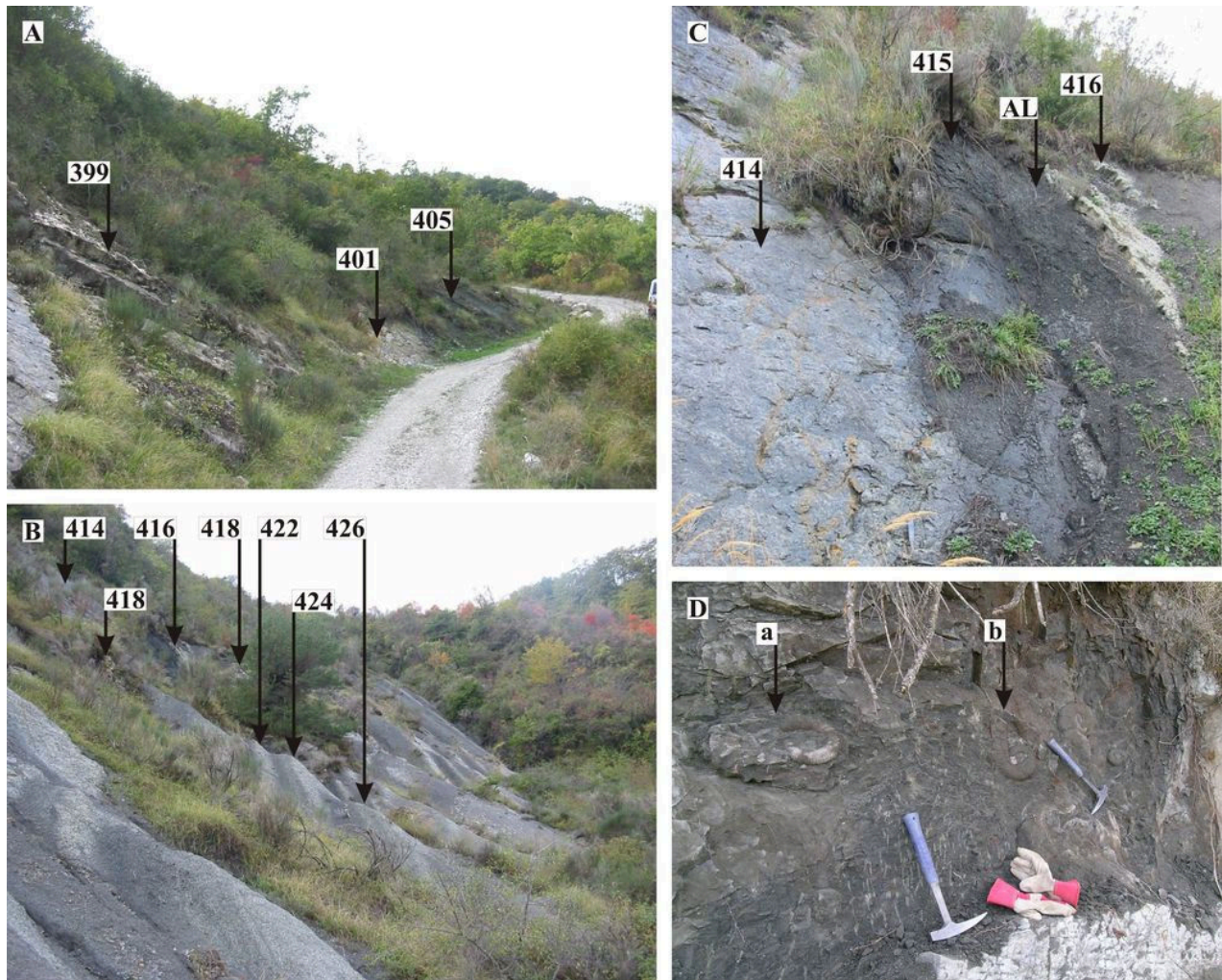


Figure 11: Field photographs of the Combe de Joinet section (CHP). Log of this section is figured Fig. 10. **A:** top of the Hauterivian-Aptian limestones and lower part of the section in 2004. **B-C:** top of the Pont de la Cerise Member and lower part of the Les Graous Member in 2004. AL: *Ammonitoceras* Level. **D:** ammonites *in situ* in the *Ammonitoceras* Level in 2009 (refigured from BERT, 2014, p. 393). a: specimen No. "28769" identified as *Caseyites esteronsensis* DELANOY *et al.*, 2018 in DELANOY *et al.* (2018, p. 58, Fig. 12.A-B), quoted as *Ammonitoceras ucetiae* DUMAS, 1876 [m] in Fig. 10. b: specimen No. AP-001, holotype of *Pseudocrioceras bournaudi* DELANOY *et al.*, 2018 in DELANOY *et al.* (2018, p. 47, Figs. 7, 8.A), quoted as *Ammonitoceras ucetiae* DUMAS, 1876 [m] in Fig. 10. Length of the hammer: 330 mm.

Apart from the *Ammonitoceras* Level, the ammonite fauna of the Pont de la Cerise Member is represented by rare and poorly preserved internal moulds (Fig. 13.D). We recognized the following taxa: *Ammonitoceras* sp., *Lithancylus grandis*, *Deshayesites* sp., *Chelonicerus* sp., *Pseudohaploceras liptoviensis* and a fragmentary Ancyloceratidae unidentifiable at the generic rank. Belemnites are represented by *Neohibolites aptiensis* (STOLLEY, 1913) and *Duvalia grasi* (DUVAL-JOUE, 1841). A level one centimetre above the *Ammonitoceras* Level yielded small internal phosphate moulds of *Macroscaphites striatisulcatus* and *Pseudohaploceras liptoviensis*.

The *Ammonitoceras* Level is a 0.20 m-thick level particularly rich in ammonites. It is located approximately 0.50 m below the upper limit of the Pont de la Cerise Member (Figs. 4, 11.C-D).

This level provided an abundant cephalopod fauna mainly consisting of deformed internal moulds of ammonites and nautiloids, and belemnite rostra (Pl. 2, figs. 1-17, 19-20). The proportion of sub-complete representatives of the heteromorphic ammonite family Ancyloceratidae, especially *Ammonitoceras*, is noteworthy (Fig. 11.D ; Pl. 2, figs. 1-5). The type section of this reference level is the Combe de Joinet section (CHP) where it crops out under particularly favourable conditions (see Chapter 5.3.4). Heteromorph ammonites are deposited without preferential orientation (Fig. 11.D). BERT (2014, p. 393) figured this level for the first time (Fig. 11.D), and DELANOY *et al.* (2018) described and figured many ammonites. We identified the following taxa in the *Ammonitoceras* Level: *Ammonitoceras ucetiae* macroconchs (Pl. 2, figs. 1-2) and microconchs (Pl. 2, figs. 3-5), *Lithancylus grandis* (Pl. 2, figs. 11-12), *Deshayesites multicostatus* (Pl.

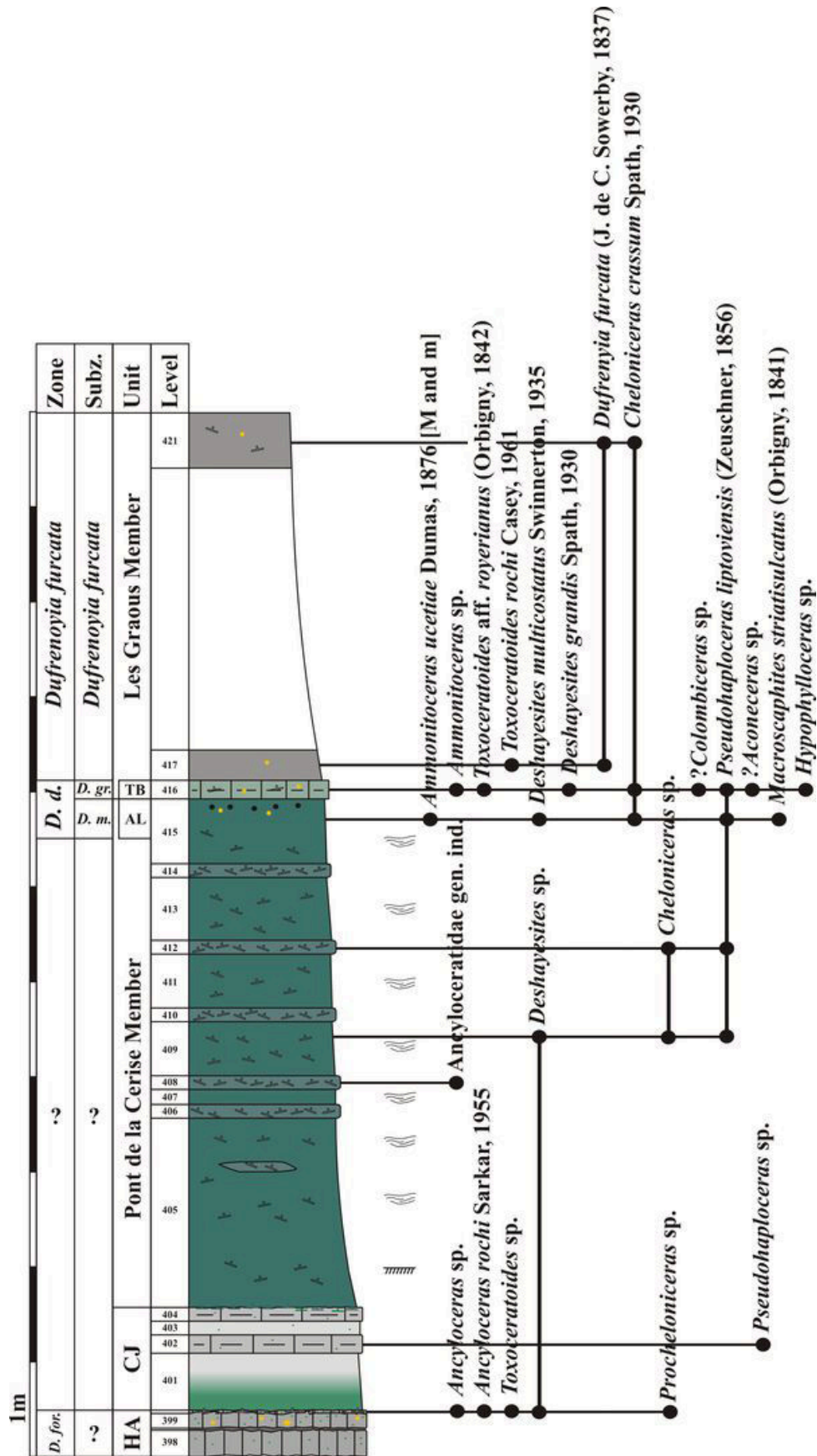


Figure 12: Log of the Pont de la Cerise section (CLE) with distribution of ammonite taxa, see Figs. 4 - 5 for caption.

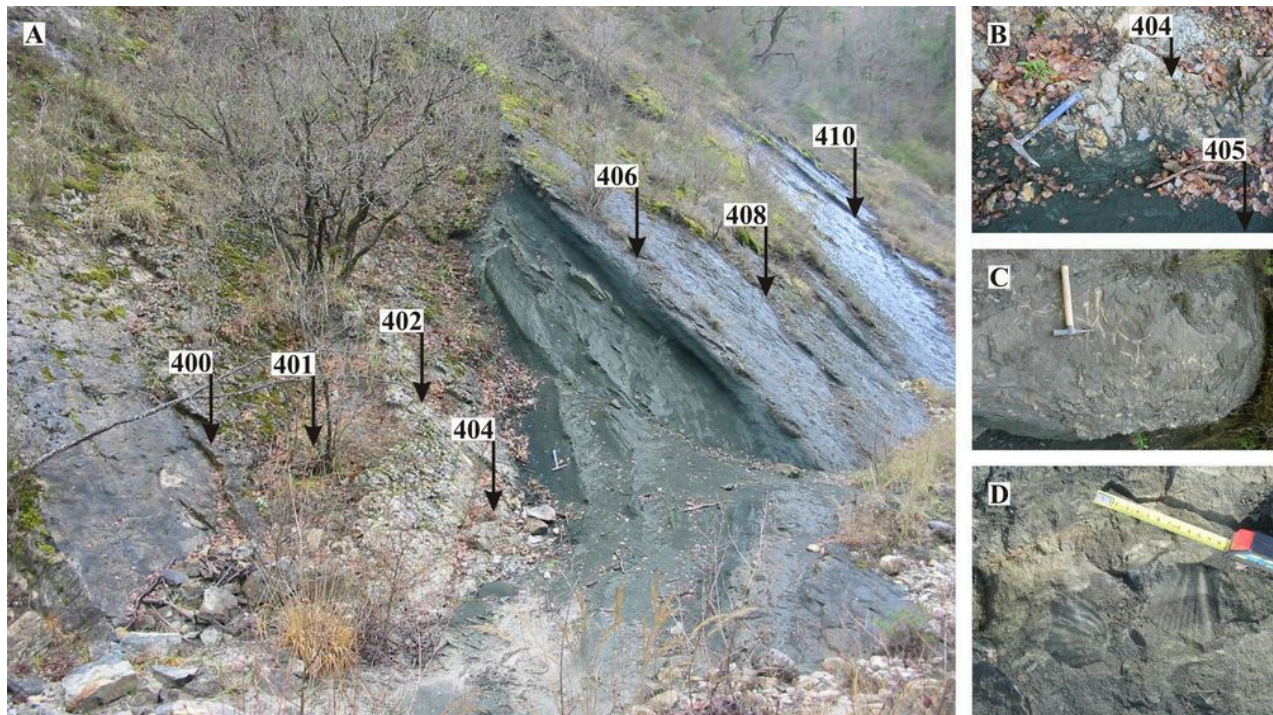


Figure 13: Field photographs of the Pont de la Cerise section (CLE). Log of this section is available in Fig. 12. **A:** top of the Hauterivian-Aptian limestones, Combe de Joinet Member and Pont de la Cerise Member in 2004. **B:** *Thalassinoides* at the top of the Combe de Joinet Member (bed 404) in 2004. Length of the hammer: 330 mm. **C:** *Thalassinoides* in Bed 408 of the Pont de la Cerise Member. Length of the hammer: 290 mm. **D:** ammonites fragments in Bed 409 (*Chelonicerases* sp. and *Pseudohaploceras* sp. in Fig. 12). Tape measure with centimetre markings.

2, figs. 6-8, 13-15, 19), *Chelonicerases crassum* (Pl. 2, figs. 16-17), *Chelonicerases* sp., *Pseudohaploceras liptoviensis* (Pl. 2, figs. 9-10), *Pseudosaynella* sp., *Macroscephites striatisulcatus* (Pl. 2, fig. 20). Belemnites are represented by *Neohibolites aptiensis* and *Duvalia grasi*. Nautiloids are frequent and represented by *Cymatoceras neckerianus*, *Eucymatoceras plicatum* (FITTON, 1836) and *Anglonautilus* sp. The level 405 of the La Valliere section (VAL) is a possible local representative of the *Ammonitoceras* Level (see Chapter 5.3.2) and additionally provided *Toxoceratoides* aff. *royerianus*. Gastropods, bivalves and ahermatypic corals are relatively rare (Pl. 2, fig. 18).

Deshayesitidae of the *Ammonitoceras* Level have a smooth siphonal band ending at a diameter between 15 and 18 mm and their ventral area is not subtabulated on the phragmocone (Pl. 2, figs. 8, 14). These characters correspond to the index species *D. multicostatus* (BERSAC & BERT, 2015). The genus *Chelonicerases* is represented by two distinct morphologies:

1. a relatively frequent one, characterized by a long *Chelonicerases* stage and a short *Procheloniceras* stage. The "intermediate stage" of *Chelonicerases parinodum* CASEY, 1961, is present in some specimens between the *Chelonicerases* and *Procheloniceras* stages (BERSAC & BERT, 2018). For these reasons,

we attributed this morphology to a primitive form of *C. crassum* (Pl. 2, figs. 16-17);

2. a much rarer second morphology, with evolute coiling and with relatively short *Chelonicerases* stage. This form we left in open nomenclature, resembles the primitive parahoplitid *Colombiceras* and could be its ancestor.

The presence of *D. multicostatus* and *C. crassum* with primitive morphology in the *Ammonitoceras* Level indicates an age corresponding to the late part of the *D. multicostatus* Subzone (*D. deshayesi* Zone) sensu BERSAC *et al.* (2012; Fig. 1). The morphology of the *Ammonitoceras* in the *Ammonitoceras* Level, with a short *Ammonitoceras* ontogenetic stage and an important spiral hiatus in macroconchs (Pl. 2, fig. 1), is consistent with such an age.

The absence of *Deshayesites* bearing a longer smooth siphonal band and a subtabulated ventral area on the phragmocone (corresponding to the index species *D. grandis*, see BERSAC & BERT, 2012), allows us to exclude the presence of the *D. grandis* Sub-zone in the *Ammonitoceras* Level.

The *Ammonitoceras* Level crops out in the following sections: Les Graous 1 (GRS1), Les Graous 2 (GRS2), La Vallière (VAL), Combe de Joinet (CHP), Pont de la Cerise (CLE) and Pont Pairé (PPR).

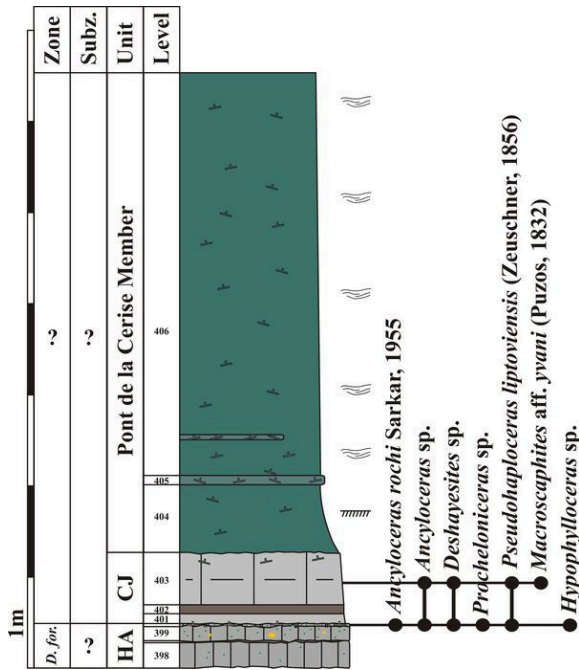


Figure 14: Log of the Combe de Marin section (CRS) with distribution of ammonite taxa, see Figs. 4 - 5 for caption.

5.2.2.3. The Les Graous Member

This member is restricted to the LFAB (Fig. 4). It starts with a marly limestone level of about 0.2 m thickness, named here the *Toxoceratoides* Bed (Figs. 4, 11.C). The type section of the Les Graous Member is at the Les Graous 1 section (GRS1).

The levels of the Les Graous Member above the *Toxoceratoides* Bed are represented in the sections at Les Graous 1 (GRS1) and Les Graous 2 (GRS2) by blue-grey sandy marls. A bundle of 6 more indurated beds is intercalated in the lower part of these marls (Figs. 6.B, 11.C-D). Within the Les Graous Member, pyrite nodules and centimetre-thick pyrite alteration beds are relatively common. The upper limit of this member can only be observed at the GRS 1 and GRS 2 sections: it corresponds to a discontinuity surface with numerous *Thalassinoides* located at the top of a 0.80 m-thick marly limestone bed (GRS 1 and GRS 2 bed 113; Figs. 5, 6.D).

The *Toxoceratoides* Bed provided a large number of deformed internal moulds of ammonites, nautiloids, belemnites (Pl. 3, figs. 7-8), bivalves, gastropods, corals, and fish scales. The ammonite fauna is represented by *Ammonitoceras* sp., *Lithancylus* sp., *Toxoceratoides* aff. *royerianus* (Pl. 3, figs. 13-16), *Deshayesites grandis* (Pl. 3, figs. 9-12), *Cheloniceras crassum* (Pl. 3, figs. 5-6), *?Colombiceras* sp., *Pseudohaploceras liptoviensis*, *?Aconeceras* sp., *Macroscaphites striatisulcatus* (Pl. 3, figs. 3-4), *Acantholytoceras* sp., *Ptyhoceras emericianum*, *Hypophylloceras* sp. (Pl. 3, fig. 1) and *Lytoceras* sp.

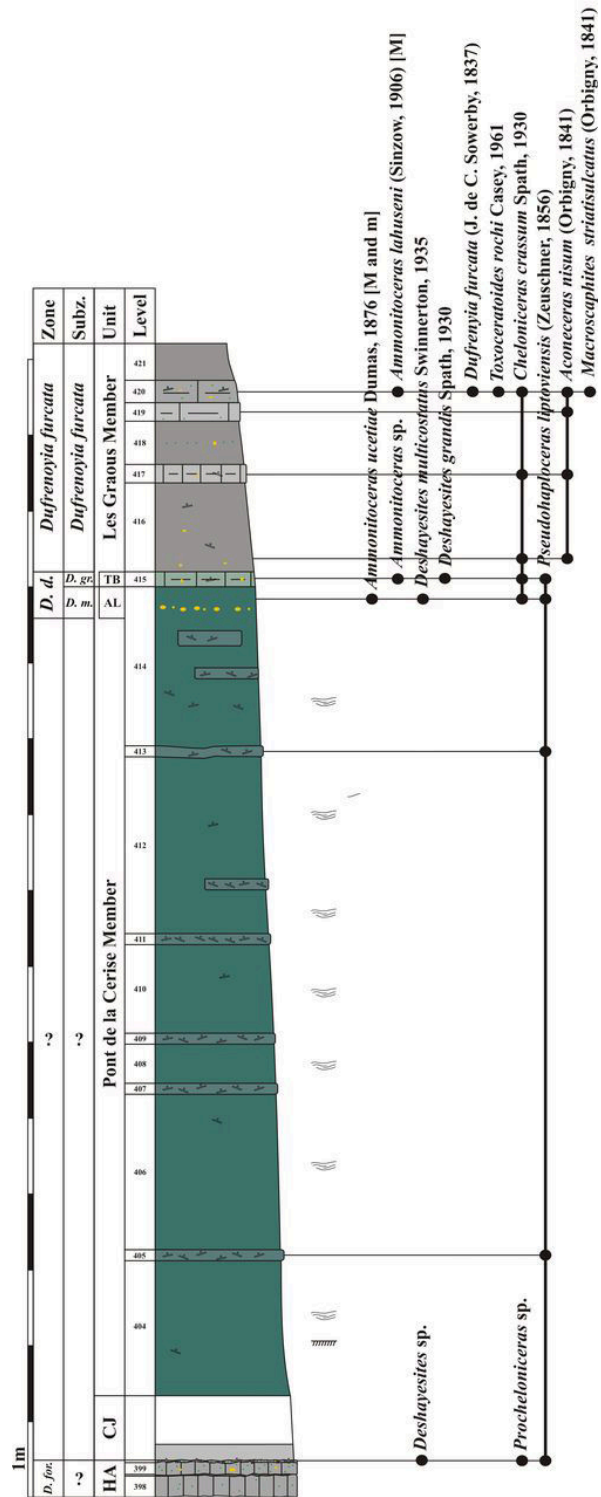


Figure 15: Log of the Pont Pairé section (PPR) with distribution of ammonite taxa, see Figs. 4 - 5 for caption.

The *Deshayesites* in this level have a well-expressed smooth siphonal band and a subtabulated ventral area on the phragmocone (Pl. 3, figs. 9, 12). The *Cheloniceras* have a long *Cheloniceras* stage and a short or absent *Procheloniceras* stage (Pl. 3, fig. 5). These characters indicate an age corresponding to the *D. grandis* Subzone (Fig. 1).



Figure 16: Field photographs of the Pont de la Cerise section (CLE). Log of this section is available in Fig. 15. **A:** Pont de la Cerise Member in 2017. **B:** top of the Pont de la Cerise Member and base of the Les Graous Member in 2017. AL: *Ammonitoceras* Level.

Above the *Toxoceratoides* Bed, the lower part of the Les Graous Member yielded an abundant ammonite fauna represented by deformed internal moulds (Fig. 6.C). Within these levels, belemnites (Pl. 5, figs. 17-21) and fish scales (Pl. 5, fig. 26) are relatively frequent. Nautiloids, echinoids (Pl. 5, figs. 27-28), brachiopods (Pl. 5, fig. 29), gastropods (Pl. 5, figs. 22-23), bivalves (Pl. 5, fig. 24) and small ahermatypic corals (Pl. 5, fig. 25) are rarer. Fossils are noticeably more frequent in the more indurated beds. Pyritic internal moulds of ammonites are extremely rare and poorly preserved. In the lower part of the grey-blue marls, we identified the following ammonitic taxa: *Ammonitoceras lahuseni* macroconchs (Pl. 4, figs. 1-3) and microconchs, *Tropaeum bowerbankii*, *Toxoceratoides rochi* (Pl. 5, figs. 1-4), *Dufrenoyia furcata* (Pl. 5, figs. 5-9), *Chelonicerus crassum* (Pl. 4, figs. 8-9), *Colombicerus crassicostratum* (Pl. 5, figs. 10-12), *Pseudohaploceras liptoviensis* (Pl. 4, figs. 6-7), *Aconecerus nisum* (Pl. 5, figs. 13-14), *Ptychoceras emericianum* (Pl. 5, fig. 15), *Macroscaphites striatisulcatus* (Pl. 5, fig. 16), *Acantholytoceras* sp. (Pl. 4, figs. 4-5), *Lytoceras* sp. (Pl. 4, figs. 10-11), *Hypophylloceras* sp. (Pl. 4, fig. 12). Nautilites are represented by *Cymatoceras neckerianus* (Pl. 4, figs. 13-14) and *Anglonautilus* sp.

In the upper part of the Les Graous Member, we identified *Toxoceratoides rochi*, *Dufrenoyia dufrenoyi* (Pl. 5, figs. 30-32), *Chelonicerus crassum*, *Pseudohaploceras liptoviensis*, *Aconecerus nisum*, *Hypophylloceras* sp.

Belemnite taxa, *Neohibolites aptiensis* (Pl. 5, figs. 17-18), *Duvalia grasi* (Pl. 5, figs. 19-20) and *Mesohibolites* sp. (Pl. 5, fig. 21), occur throughout the Les Graous Member.

The grey-blue marls of the Les Graous Member are assigned to the *D. furcata* Zone (Fig. 1): their lower part is assigned to the *D. furcata* Subzone and their upper part to the *D. dufrenoyi* Subzone.

5.3. DESCRIPTIONS OF THE STUDIED SECTIONS

Eight sections were sampled between 2000 and 2017. They are located along an E-W axis on either side of the village of Les Ferres (Fig. 3). These sections are (from W to E):

5.3.1. Les Graous sections (GRS1 and GRS2)

These sections are located about 1.65 km to the WSW of the village of Les Ferres. Les Graous 1 (GRS 1, Fig. 3, lat.: 43°50'31.8"N, long.: 7°4'35.4"E) crops out on the eastern flank of the Les Graous cuesta and Les Graous 2 (GRS 2, Fig. 3, lat.: 43°50'31.0"N, long.: 7°4'27.3"E) on its western flank. Because of their similar lithostratigraphy, the log of GRS 1 only is figured (Fig. 5). Field photographs and log section of GRS 1 were given by BERT & BERSAC (2011, Figs. 30, 32-37, 40-47), but the exact location of this section was not indicated because fieldwork had not been completed then. In this section, the marly Aptian can be observed continuously from the top of the Hauterivian-Aptian limestones to the Albian (Fig. 6.B, .D). The Combe de Joinet Member is absent and the Pont de la Cerise Member lies directly on the Hauterivian-Aptian limestones. The Pont de la Cerise Member is only 0.90 m thick in GRS 1 (Figs. 5, 6.A). A bed of phosphate nodules can be observed at 15 cm above its base. At the top of this member the *Ammonitoceras* Level provided some poorly preserved ammonite fragments. In addition to the ammonite taxa cited in Fig. 5, the Pont de la Cerise Member provided *Cymatoceras neckerianus* (20 cm above its base), *Anglonautilus* sp. (30 cm above its base) and several specimens of *Neohibolites aptiensis*.



The *Toxoceratoides* Bed is present at the base of Les Graous Member (Figs. 5, 6.A). It is relatively marly here and yielded a fairly abundant fauna of ammonites, belemnites and nautiloids. Above the *Toxoceratoides* Bed, the first 5 metres of this member are represented by a bundle of 6 sandy and marly limestone beds separated from each other by marly intervals (Figs. 5, 6.B). A rich ammonite fauna assigned to the *D. furcata* Subzone (Fig. 6.C ; Pl. 4, figs. 8-12; Pl. 5, figs. 1-16) was collected mainly from the first 5 sandy and marly limestone beds and Bed 99 (Fig. 5).

The upper part of the Les Graous Member is attributed to the *D. dufrenoyi* Subzone. It is marlier and only the sandy and marly limestone beds yielded ammonites. Bed 110 yielded the earliest *D. dufrenoyi* (Fig. 5). The upper limit of this member is represented by a sandy limestone bed (Bed 113, Fig. 5) that is relatively rich in ammonites (especially in *D. dufrenoyi*, Pl. 5, fig. 32). It is overlain by upper Aptian dark glauconitic marls. All the material from "les Graoux" published by DELANOY *et al.* (2018) was collected from Beds 99, 100 and 102 of the GRS 1 section in our presence.

5.3.2. La Vallière section (VAL)

RAGAZZI (1982, p. 73) and BRÉHERET (1997, p. 267, Figs. 114-115) studied a section located in the La Vallière area (Fig. 3, lat.: 43°50'32.6"N, long.: 7°5'29.9"E). This section is nowadays almost completely covered by vegetation and located in a private fenced property.

We studied an outcrop near this section (VAL section, Figs. 3, 7, lat.: 43°50'32.1"N, long.: 7°5'48.4"E). It is located in a half-graben, against a vertical decametric palaeo-escarpment, about 0.6 km SSW of the village of Les Ferres (Fig. 8.A). A cross-section of this tectonic structure was figured by DARDEAU & GRACIANSKY (1987, Figs. 1-2). In this outcrop, the Pont de la Cerise Member reaches its maximum thickness (15 m) and is overlain by a 0.2 m-thick sandstone level (Bed 417, Fig. 7), which is the probable local equivalent of the *Toxoceratoides* Bed (Fig. 8.A). Bed 405, located at 3.70 m from the base of the Pont de la Cerise Member (Figs. 7, 8.B), yielded a relatively abundant, but fragmentary, ammonite fauna very similar to that of the *Ammonitoceras* Level. A level located at the top of the Pont de la Cerise Member (top of Level 416, Fig. 7) yielded rare representatives of *Toxoceratoides*, *Deshayesites* and *Chelonoceras*, but they are too fragmentary to be identified to specific level. Nevertheless, the lithological and palaeontological characteristics of this last bed are not equivalent to the *Ammonitoceras* Level. For this reason, we consider Bed 405 as a local equivalent of the *Ammonitoceras* Level. This all implies that the glauconitic sandstones over-

lying the *Ammonitoceras* Level of VAL are therefore considerably more developed here than in other sections (Fig. 17). The Les Graous Member of VAL is almost completely covered by vegetation and thus could not be studied.

5.3.3. La Graou section (GRO)

The La Graou section (GRO, lat.: 43°50'36.2"N, long.: 7°6'0.0"E) was described by RAGAZZI (1982, p. 73), then by BRÉHERET (1997, p. 264, Fig. 114). It is located 0.5 km SW to the village of Les Ferres (Figs. 3, 9). The Lower Aptian is nowadays almost totally covered by vegetation and only a small part of the Les Graous Member can be observed.

5.3.4. The Combe de Joinet section (CHP)

The CHP section (lat.: 43°50'42.7"N, long.: 7°6'9.5"E) is located 0.8 km ESE of the village of Les Ferres, along a track leading to the Pont de la Cerise (Figs. 3, 10 - 11). BERT (2014, p. 393) figured the *Ammonitoceras* Level at the CHP section for the first time (Fig. 11.D). DELANOY *et al.* (2018, Fig. 2) published a log of this section and described and figured many ammonites, mainly from the *Ammonitoceras* Level (Bed 107 in DELANOY *et al.*, 2018 = Bed 415 in the present work, Figs. 10, 11.C-D). Unfortunately, several levels located at the base of the marly Aptian were not figured in the log of DELANOY *et al.* (2018, Fig. 2), including the Combe de Joinet Member. For this reason, we did not use the bed numbering system of this latter work for our own log (Fig. 10). At CHP, the Combe de Joinet Member (Bed 401) crops out in good conditions along the track between the D1 road and the Pont de la Cerise (Figs. 3, 11.A). It is represented by massive sandy and marly limestones and provided some rare fragments of ammonite. The *Ammonitoceras* Level crops out under good conditions, allowing the collecting of abundant ammonites (DELANOY *et al.*, 2018; Fig. 11.C-D ; Pl. 2, figs. 1-19). Only the lower part of the Les Graous Member is visible, its upper part being covered with dense vegetation (Fig. 11.B).

5.3.5. The Pont de la Cerise section (CLE)

This section (lat.: 43°50'42.7"N, long.: 7°6'9.5"E) was first published by RAGAZZI (1982, p. 75) who named it "Pont de la Cerise" despite the fact that another section closer to the Pont de la Cerise locality was named "Combe de Marin" by RAGAZZI (1982, p. 77 and Fig. 3; Fig. 3 and see Chapter 5.3.6). The CLE section crops out on the right bank of the Esteron River in the locality of the Clot d'Estéron, 1.95 km east of the village of Les Ferres (Figs. 3, 12). The top of the Hauterivian-Aptian limestones, the Combe de Joinet and Pont de la Cerise members, as well as the *Toxoceratoides* Bed crop out under favourable conditions (Fig. 13.A). As RAGAZZI noticed (1982, p. 76), the Combe de Joinet Member (Beds 401 to 404 in CHP) is represented by 1.1 m of marly limestone and two decimetric limestone beds (Figs. 12, 13.A). The top of this member is intensely bioturbated (Fig. 13.B). The Pont de la Cerise Member crops out in very

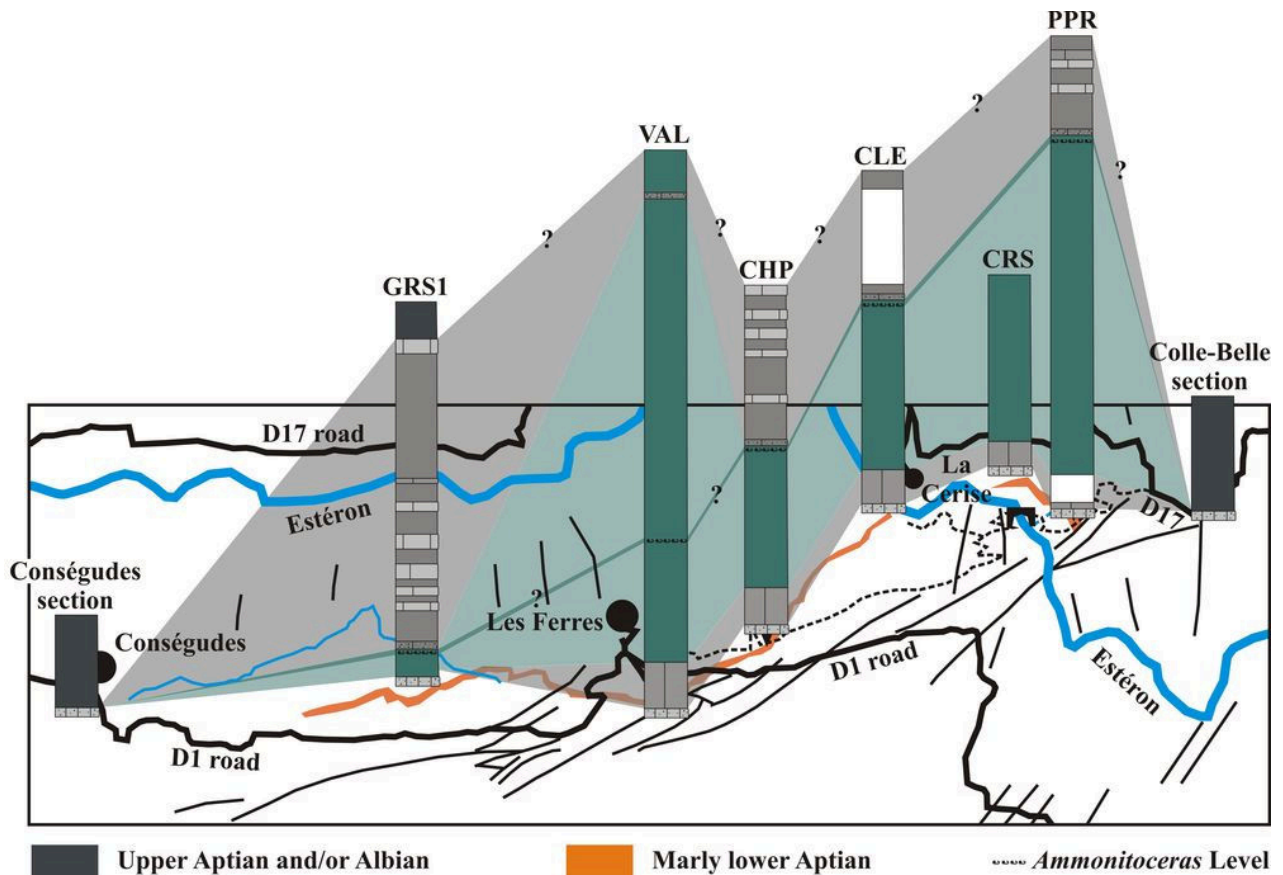


Figure 17: Lithostratigraphic correlation between the studied sections and the adjacent sections.

good conditions and is 5.5 m thick at that point (Fig. 13.A, .C-D). In its lower and middle parts, rare and poorly preserved fragments of *Ancyloceras* sp., *Deshayesites* sp., *Pseudohaploceras liptoviensis* and *Macroscaphites* aff. *yvani*.

only the *Toxoceratoides* Bed (Bed 416, Fig. 12) of the Les Graous Member crops out under good conditions, the rest of the formation being entirely covered by vegetation.

5.3.6. The Combe de Marin section (CRS)

This section was figured by RAGAZZI (1982, p. 77 and Fig. 3). It is located on the left bank of the Estéron River, opposite to the Pont de la Cerise locality, 2.69 km east of the village of Les Ferres (Fig. 3). The CRS section crops out at the top of a cliff mainly formed by the Hauterivian-Aptian limestones (Fig. 14). Our own observations differ slightly from those of RAGAZZI (1982): only the Combe de Joinet Member and the lower 5.4 m of the Pont de la Cerise Member are currently observable. The rest of the Aptian marls is covered by vegetation. The Combe de Joinet Member is represented by 0.80 m of marly limestones, with a dark brown clay level, 0.10 m thick, in its lower part. The ammonite fauna is rare and fragmen-

tary, but includes: *Ancyloceras* sp., *Deshayesites* sp., *Pseudohaploceras liptoviensis* and *Macroscaphites* aff. *yvani*.

5.3.7. The Pont Pairé section (PPR)

RAGAZZI (1982) already mentioned this section ("ravin de Pont de Payre" in RAGAZZI, 1982, p. 77) where she observed that the Pont de la Cerise Member was reduced to a thickness of 3.5 metres. This section (lat.: 43°51'11.6"N, long.: 7°7'49.0"E, Fig. 15) crops out in a ravine perpendicular to the track between the D17 road and Pont de la Cerise (Fig. 3). Our own observations differ from RAGAZZI's (1982): the Pont de la Cerise Member is separated from the Hauterivian-Aptian limestones by a metre-thick covered section and is 11 m-thick (Figs. 15, 16.A-B). This member can be observed under good conditions along the ravine, for a distance of about 100 m. The *Ammonitoceras* Level (top of Bed 414 in PPR, see Fig. 15) is present at its top (Fig. 16.B), immediately above a level particularly rich in large pyrite nodules. The levels of the Pont de la Cerise Member below the *Ammonitoceras* Level reach their maximum thickness in this section. Ammonites are very rare in the Pont de la Cerise Member in this section, except in the *Ammonitoceras* Level. Only the first 3.3 m of the Les Graous Member are observable (Fig. 16.B); they are richer in glaucony than in the other sections. The *Toxoceratoides* Bed is marlier than in the CHP and CLE sections, and yielded some ammonites (Figs. 15, 16.B).

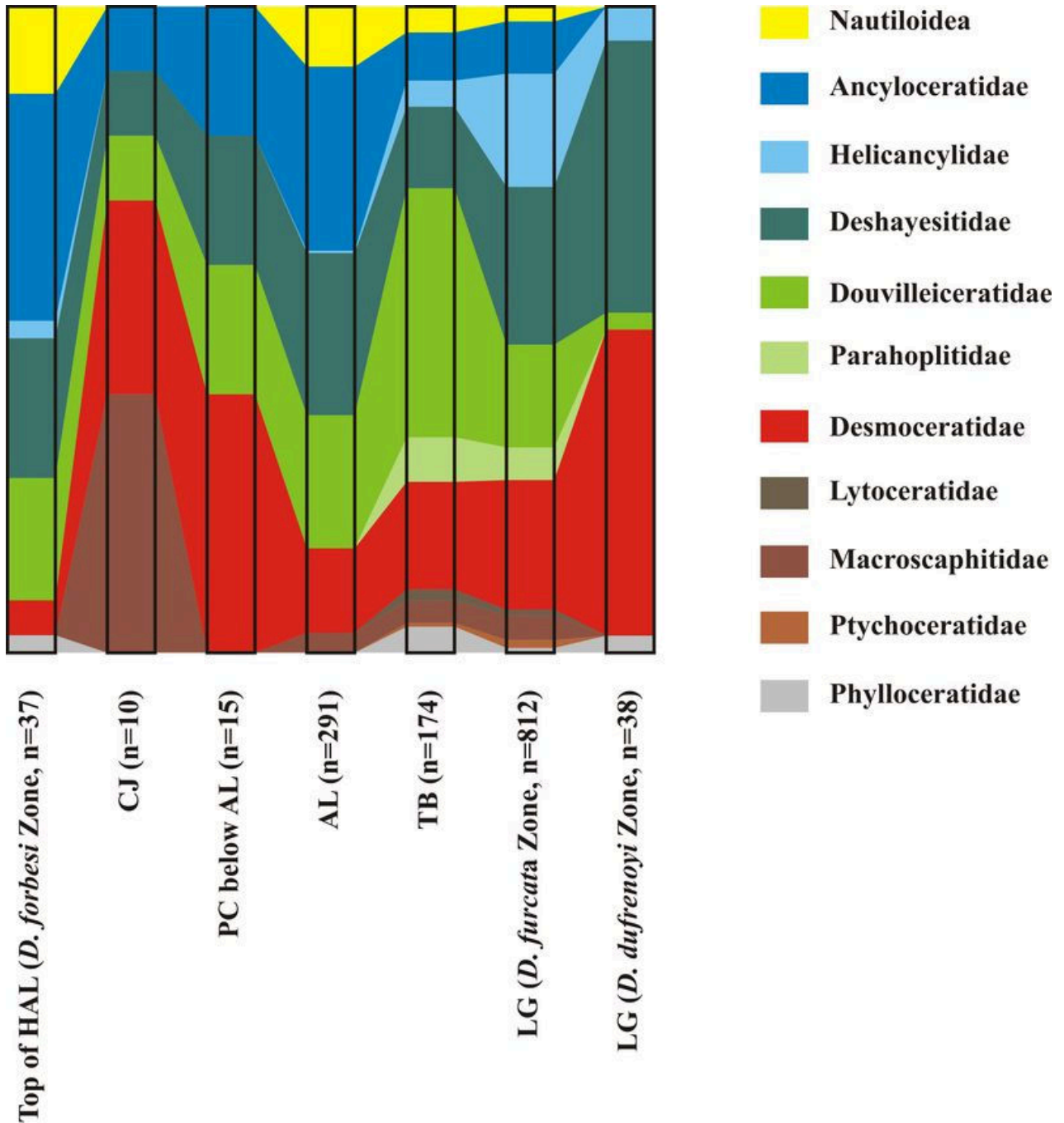


Figure 18: Nautiloidea and Ammonoidea faunal spectrum of the lower Aptian of les Ferres area. HAL: Hauterivian-Aptian limestones, n: number of specimens studied, CJ: Combe de Joinet Member, PC: Pont de la Cerise Member, AL: *Ammonitoceras* Level, TB: *Toxoceratoides* Bed, LG: Les Graous Member.

5.4. LATERAL VARIATIONS OF THE MARLY APTIAN SERIES

The Combe de Joinet Member appears at VAL, where it reaches its maximum thickness (taking into account that this member could not be measured at PPR). It gets thinner towards the east (Fig. 17). The Pont de la Cerise Member appears at GRS 1 and GRS 2 with a thickness of only 0.9 m. The interval of this member beneath the *Ammonitoceras* Level gradually thickens towards the east (taking into account

that we have no thickness data at CRS, Fig. 17) reaching their maximum thickness at PPR. The interval of the Pont de la Cerise Member above the *Ammonitoceras* Level are generally a few centimeters-thick, except at VAL where they exceptionally reach a thickness of 11 m (Figs. 7, 8.A, 17). This may be due to the palaeogeographic context of VAL: the presence of a decametric fault palaeo-escarpment in this area probably allowed an accumulation of a large amount of sediment above the *Ammonitoceras* Level (DARDEAU & GRACIANSKY, 1987). Our observations differ from those of RA-



GAZZI (1982) who observed a thickness reduction of the Pont de la Cerise Member ("accumulation glauconitique" in RAGAZZI, 1982, p. 78) at CRS and PPR (Fig. 17) compared to more westerly sections. The thickness variation of the Les Graous Member could not be studied because this latter is completely visible only at GRS 1 and GRS 2.

The nearest outcrop west of GRS 2 (Fig. 3) confirms the westward thickness reduction of the Pont de la Cerise Member, as already observed by RAGAZZI (1982). This section, named "Conségudes" by RAGAZZI (1982, Fig. 3), crops out along the D1 road next to the village of Conségudes (lat.: 43°50'26.5"N, long.: 7°2'58.3"E, Fig. 3): the "Early Aptian unconformity" (COTILLON, 2010, p. 4) is located at the top of the upper Barremian limestones and is directly covered by Albian dark marlstones (RAGAZZI, 1982, p. 73 and Pl. 3; Fig. 17). The Aptian is therefore totally absent in this section.

East of the study area, the Aptian is also totally missing in the Colle-Belle section (COTILLON, 1971, p. 147, lat.: 43°51'10.4"N, long.: 7°8'32.3"E, Fig. 17), with Albian marlstones lying directly on Barremian limestones (COTILLON, 1971, p. 132 and 147; Fig. 17). The sudden demise of the marly Aptian between the PPR and the Colle-Belle section is probably related to the presence of a decametric synsedimentary fault forming the eastern limit of the LFAB.

North of the study area, Aptian deposits can be observed only on the northern flanks of the Roccaforte and Vial mountains and near the village of Tournefort, where they are represented by plurimetric dark marlstones with pyritized ammonites characteristic of the pelagic domain (RAGAZZI, 1982, p. 61 and personal observations).

To the south, the Aptian is missing or represented in the Coursegoules and Ruth Mountain areas by thin reworked layers (COTILLON, 1971, p. 145; BERSAC *et al.*, 2010; BERT & BERSAC, 2011, Figs. 4-6, 20, 67-70; BERT *et al.*, 2013, p. 361, Fig. 9).

5.5. AMMONITE AND NAUTILOID FAUNAL SPECTRUM

The faunal spectrum study of the ectocochleate cephalopods of the LFAB is based on 1,262 ammonites and 58 nautiloids collected or studied *in situ*, and by 29 ammonites and 1 nautiloid published in DELANOY *et al.* (2018).

The most represented ammonites families in the lower Aptian of the LFAB are Ancyloceratidae, Deshayesitidae, Douvilleiceratidae, and Desmoceratidae (Fig. 18). This faunal spectrum is consistent with the palaeogeography of the LFAB during the early Aptian times, located in the outer neritic domain. Ancyloceratidae and Nautiloidea (WARD, 1987; HEWITT, 1989; WESTERMANN, 1990, 1996) have a preferential tro-

pism for the neritic domain, Deshayesitidae are open oceanic surface organisms, Douvilleiceratidae are opportunistic (BULOT *et al.*, 2005; DUTOUR, 2005) and Desmoceratidae, Phylloceratidae and Lytoceratidae are deeper water environmental representatives (WESTERMANN, 1990, 1996; BULOT *et al.*, 2005).

The most spectacular character of the lower Aptian ammonite fauna of the LFAB is the abundance of the heteromorphic genus *Ammonitoceras* in the *D. deshayesi* and *D. furcata* Zones, with 80.2% of the Ancyloceratidae of the *D. deshayesi* Zone and 91.9% for the *D. furcata* Zone respectively. Such an abundance has never been reported previously in the literature (KAKABADZE, 1981; KAKABADZE & HOEDEMAEKER, 2004). It may be related to particularly favourable environmental conditions in the LFAB.

The scarcity of representatives of the genus *Trochaeum* in the LFAB may be related to the fact that this is a boreal taxon (CASEY, 1960; AGUIRRE-URRETA, 1985). Nautiloidea never represent more than 13.5% of the ectocochleate cephalopod fauna (Fig. 18).

The proportion of each group of ectocochleate cephalopods varies over the stratigraphic interval considered: excluding data from the Combe de Joinet Member and the Pont de la Cerise Member below the *Ammonitoceras* Level due to the small amount of material collected (respectively 10 and 15 specimens; Fig. 18), there is a decrease in the proportion of Nautiloidea and Ancyloceratidae over time and an increase of Deshayesidae and Desmoceratidae (the latter being related to the "bloom" of the genus *Aconeceras* HYATT, 1903, in the *D. furcata* Zone). This is strongly consistent with the transgressive trend of the studied series (see Chapter 2), where open marine environment taxa become progressively more prevalent. At the top of the studied series, in the *D. dufrenoyi* Subzone, this phenomenon leads to a faunal spectrum similar to that observed in hemipelagic environments, where the ammonite populations are mainly composed of Deshayesitidae and Desmoceratidae (DUTOUR, 2005, Fig. 22). The proportion of Helicancyllidae is higher when Ancyloceratidae decrease, as a result of ecological interactions between these two families.

6. Conclusion

The marly deposits of the Les Ferres Aptian Basin (LFAB) are located in one of the few areas of the southern margin of the Vocontian Basin where ammonites from the *D. deshayesi* – *D. furcata* Zones can be studied under favourable conditions. An abundant bed-by-bed collection of ammonites, constrained at the subzone level and represented by a significant proportion of adult or subadult specimens was collected. The lithostratigraphy and biostratigraphy of the marly Aptian of Les Ferres presented in the present article aims to serve as a



basis for our future palaeontological work in this area. The whole marly Aptian of the LFAB is represented by a newly described lithological unit: the Les Graous Formation. This unit is divided into 3 members (from bottom to top): 1) the Combe de Joinet, 2) the Pont de la Cerise and 3) the Les Graous Members. Most of the collected ammonites come from a precise level at the top of the Pont de la Cerise Member, the *Ammonitoceras* Level (upper *D. multicostatus* Subzone, *D. deshayesi* Zone), and from the lower part of the Les Graous Member (*D. grandis* and *D. furcata* Subzones). One of the most remarkable features of this fauna is the abundance of the heteromorphic ammonite genus *Ammonitoceras* (Ancyloceratidae). The abundant collected material allows us to formulate hypotheses on the intraspecific variability and the evolutionary modalities of several taxons of the families Ancyloceratidae, Douvilleiceratidae, Deshayesitidae, Parahoplitidae, and Desmooceratidae, which will be the topic of forthcoming contributions.

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Plates

Plate 1: Cephalopods of the last bed of the Hauterivian-Aptian limestones (*Deshayesites forbesi* Zone). All specimens: BERSAC's collection.

Fig. 1: *Ancyloceras* sp., SBC.06061-00007/CLE037, CLE section, Bed 400.

Fig. 2: *Ancyloceras rochi* SARKAR, 1955, SBC.06061-00007/CLE032, CLE section, Bed 400.

Fig. 3: *Ancyloceras* sp., SBC.06061-00008/CRS022, CRS section, Bed 400.

Figs. 4-5: *Deshayesites* sp., SBC.06061-00007/CLE030, CLE section, Bed 400.

Fig. 6-7: *Procheloniceras* sp., SBC.06061-00007/CLE039, CLE section, Bed 400.

Fig. 8: *Pseudohaploceras* sp., SBC.06061-00008/CRS023, CRS section, Bed 400.

Fig. 9: *Cymatoceras neckerianus* (PICTET, 1847), SBC.06061-00008/CRS014, CRS section, Bed 400.

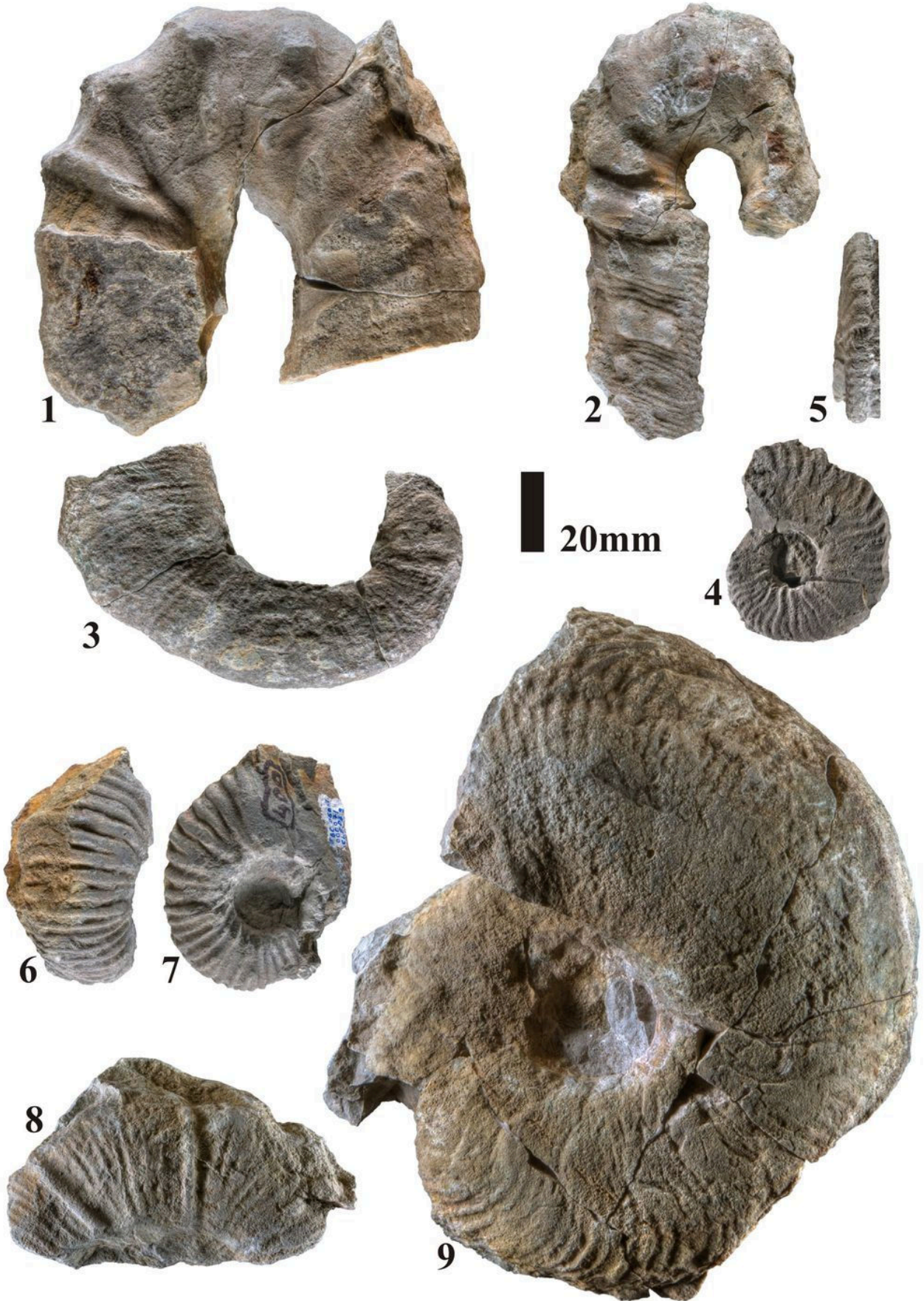




Plate 2: Fossils of the *Ammonitoceras* Level (top of the Pont de la Cerise Member, *Dehayesites deshayesi* Zone, *Dehayesites multicostatus* Subzone). Scale bar 50 mm for figures 1 to 17. Scale bar 10 mm for figures 18 to 20. All specimens: BERSAC's collection.

Figs. 1-2: *Ammonitoceras ucetiae* DUMAS, 1876, macroconch, SBC.06061-00006/CHP249, CHP section, Bed 415.

Figs. 3-5: *Ammonitoceras ucetiae* DUMAS, 1876, microconch, SBC.06061-00006/CHP276, CHP section, Bed 415.

Figs. 6-8: *Dehayesites multicostatus* SWINNERTON, 1935, SBC.06061-00006/CHP168, CHP section, Bed 415.

Figs. 9-10: *Pseudohaploceras liptoviensis* (ZEUSCHNER, 1856), SBC.06061-00006/CHP274, CHP section, Bed 415.

Figs. 11-12: *Lithancylus grandis* (J. de C. SOWERBY, 1828), SBC.06061-00006/CHP275, CHP section, Bed 415.

Figs. 13-15: *Dehayesites multicostatus* SWINNERTON, 1935, SBC.06061-00006/CHP241, CHP section, Bed 415.

Figs. 16-17: *Cheloniceras crassum* SPATH, 1930, SBC.06061-00006/CHP171, CHP section, Bed 415.

Fig. 18: Unidentified ahermatypic coral. SBC.06061-00006/CHP455, CHP section, Bed 415.

Fig. 19: *Dehayesites multicostatus* SWINNERTON, 1935, SBC.06061-00006/CHP072, CHP section, Bed 415.

Fig. 20: *Macroscaophites striatisulcatus* (ORBIGNY, 1841), microconch, SBC.06061-00007/CLE015, CLE section, Bed 415.





Plate 3: Cephalopods of the *Toxoceratoides* Bed (base of the Les Graous Member, *Dehayesites deshayesi* Zone, *Dehayesites grandis* Subzone). All specimens: BERSAC's collection.

Figs. 1-2: *Pylloceras* (*Hypophylloceras*) sp., SBC.06061-00007/CLE005, CLE section, Bed 416.

Figs. 3-4: *Macroscaphites striatisulcatus* (ORBIGNY, 1841), macroconch, SBC.06061-00006/CHP279, CHP section, Bed 416.

Figs. 5-6: *Chelonicerias crassum* SPATH, 1930, SBC.06061-00006/CHP055, CHP section, Bed 416.

Figs. 7-8: *Anglonautilus* sp., SBC.06061-00006/CHP278, CHP section, Bed 416.

Figs. 9-10: *Dehayesites grandis* SPATH, 1930, SBC.06061-00001/GRS350, GRS section, Bed 97.

Figs. 11-12: *Dehayesites grandis* SPATH, 1930, SBC.06061-00006/CHP124, CHP section, Bed 416.

Figs. 13-14: *Toxoceratoides* aff. *royerianus* (ORBIGNY, 1842), SBC.06061-00006/CHP128, CHP section, Bed 416.

Figs. 15-16: *Toxoceratoides* aff. *royerianus* (ORBIGNY, 1842), SBC.06061-00006/CHP100, CHP section, Bed 416.





Plate 4: Cephalopods of the Les Graous Member (*Dufrenoyia furcata* Zone, *Dufrenoyia furcata* Subzone). All specimens: BERSAC's collection.

Figs. 1-3: *Ammonitoceras lahuseni* (SINZOW, 1906), macroconch, SBC.06061-00006/CHP277, CHP section, Bed 419.

Figs. 4-5: *Acantholytloceras* sp., SBC.06061-00006/CHP273, CHP section, Bed 417.

Figs. 6-7: *Pseudohaploceras liptoviensis* (ZEUSCHNER, 1856), SBC.06061-00006/CHP074, CHP section, Bed 419.

Figs. 8-9: *Cheloniceras crassum* SPATH, 1930, SBC.06061-00001/GRS341, GRS 1 section, Bed 100.

Figs. 10-11: *Lytoceras* sp., SBC.06061-00001/GRS210, GRS 1 section, Bed 102.

Fig. 12: *Pylloceras* (*Hypophylloceras*) sp., SBC.06061-00001/GRS146, GRS 1 section, Bed 100.

Figs. 13-14: *Cymatoceras neckerianus* (PICTET, 1847), SBC.06061-00001/GRS211, GRS 1 section, Bed 102.





Plate 5: Fossils of the Les Graous Member. Figures 1-29: *Dufrenoyia furcata* Zone, *Dufrenoyia furcata* Subzone. Figures 30-32: *Dufrenoyia furcata* Zone, *Dufrenoyia dufrenoyi* Subzone. All specimens: BERSAC's collection.

Figs. 1-2: *Toxoceratoides rochi* CASEY, 1961, SBC.06061-00001/GRS199, GRS 1 section, Bed 100.
Figs. 3-4: *Toxoceratoides rochi* CASEY, 1961, SBC.06061-00001/GRS555, GRS 1 section, Bed 102.
Fig. 5: *Dufrenoyia furcata* (J. de C. SOWERBY, 1837), SBC.06061-00001/GRS073, GRS 1 section, Bed 100.
Figs. 6-7: *Dufrenoyia furcata* (J. de C. SOWERBY, 1837), SBC.06061-00001/GRS683, GRS 1 section, Bed 102.
Fig. 8: *Dufrenoyia furcata* (J. de C. SOWERBY, 1837), SBC.06061-00001/GRS566, GRS 1 section, Bed 102.
Fig. 9: *Dufrenoyia furcata* (J. de C. SOWERBY, 1837), SBC.06061-00001/GRS597, GRS 1 section, Bed 105.
Fig. 10: *Colombiceras crassicostatum* (ORBIGNY, 1841), SBC.06061-00001/GRS282, GRS 1 section, Bed 102.
Figs. 11-12: *Colombiceras crassicostatum* (ORBIGNY, 1841), SBC.06061-00001/GRS394, GRS 1 section, Bed 99.
Fig. 13: *Aconeceras nisum* (ORBIGNY, 1841), SBC.06061-00001/GRS175, GRS 1 section, Bed 100.
Fig. 14: *Aconeceras nisum* (ORBIGNY, 1841), SBC.06061-00001/GRS685, GRS 1 section, Bed 102.
Fig. 15: *Ptychoceras emericianum* ORBIGNY, 1842, SBC.06061-00001/GRS518, GRS 1 section, Bed 102.
Fig. 16: *Macroscaphites striatusulcatus* (ORBIGNY, 1841), microconch, SBC.06061-00001/GRS482, GRS 1 section, Bed 104.
Fig. 17: *Neohibolites aptiensis* (STOLLEY, 1913), SBC.06061-00001/GRS677, GRS 1 section, Bed 101.
Fig. 18: *Neohibolites aptiensis* (STOLLEY, 1913), SBC.06061-00001/GRS682, GRS 1 section, Bed 102.
Fig. 19: *Duvalia grasi* (Duval-Jouve, 1841), SBC.06061-00001/GRS468, GRS 1 section, Bed 104.
Fig. 20: *Duvalia grasi* (Duval-Jouve, 1841), SBC.06061-00001/GRS674, GRS 1 section, Bed 102.
Fig. 21: *Mesohibolites* sp., SBC.06061-00001/GRS679, GRS 1 section, Bed 102.
Fig. 22: *Semisolarium* sp., SBC.06061-00001/GRS295, GRS 1 section, Bed 102.
Fig. 23: *Ceratosiphon* sp., SBC.06061-00001/GRS673, GRS 1 section, Bed 102.
Fig. 24: *Aucellina aptiensis* (ORBIGNY, 1850), SBC.06061-00001/GRS491, GRS 1 section, Bed 99.
Fig. 25: Unidentified ahermatypic coral, SBC.06061-00006/CHP228, CHP section, Bed 424.
Fig. 26: Cycloid fish scale, SBC.06061-00001/GRS369, GRS 1 section, Bed 99.
Fig. 27: *Toxaster collegnoi* SISMONDA, 1843, SBC.06061-00001/GRS675, GRS 1 section, Bed 106.
Fig. 28: *Toxaster collegnoi* SISMONDA, 1843, SBC.06061-00001/GRS676, GRS 1 section, Bed 99.
Fig. 29: ?*Cyrtothyris* sp., SBC.06061-00001/GRS670, GRS 1 section, Bed 102.
Figs. 30-31: *Dufrenoyia dufrenoyi* (ORBIGNY, 1841), SBC.06061-00001/GRS685, GRS 1 section, Bed 110.
Fig. 32: *Dufrenoyia dufrenoyi* (ORBIGNY, 1841), SBC.06061-00001/GRS592, GRS 1 section, Bed 113.

