

## Carnets Geol. 24 (3)

E-ISSN 1634-0744 DOI: 10.2110/carnets.2024.2403

## The Kalkowsky Project - Chapter V Asymmetric ooids from the Yacoraite Formation (Argentina)

## Bruno R.C. GRANIER<sup>1, 2</sup> Philippe LAPOINTE<sup>3</sup>

**Abstract:** Asymmetric ooids are documented in a brackish Maastrichtian to Danian paleolake in NW Argentina. Their distinctive asymmetric growth pattern is likely related to an uneven distribution of the Extracellular Polymeric Substances (EPS) around the coated allochem, within which calcite fibers (*i.e.*, the 'fibrite') have grown. This pattern is unlikely to be mistaken for that of other 'eccentric' ooids, such as wobbly ooids, spiny ooids, hiatus ooids, half-moon ooids, 'broken' ooids *sensu lato*, or collapsed oo-molds (referred to as 'distorted' ooids).

#### **Keywords:**

- ooid;
- microbial carbonates;
- Salta;
- Argentina;
- Maastrichtian-Danian

**Citation:** GRANIER B.R.C. & LAPOINTE Ph. (2024).- The KALKOWSKY Project - Chapter V. Asymmetric ooids from the Yacoraite Formation (Argentina).- *Carnets Geol.*, Madrid, vol. 24, no. 3, p. 75-82. DOI: 10.2110/carnets.2024.2403

**Résumé :** *Le Projet KALKOWSKY - Chapitre V. Ooïdes asymétriques de la Formation Yacoraïte (Argentine).-* Nous décrivons des ooïdes asymétriques provenant d'un paléolac saumâtre d'âge Maastrichtien à Danien du nord-ouest de l'Argentine. Leur mode de croissance particulier, *i.e.*, asymétrique, est probablement lié à une répartition inégale des substances polymériques extracellulaires (SPE) autour du grain cortiqué, au sein desquelles les fibres de calcite (*i.e.*, la "fibrite") se sont développées. Il est peu probable que ce mode de croissance puisse être confondu avec celui d'autres ooïdes "excentriques", tels que les ooïdes bancals, les ooïdes épineux, les ooïdes hiataux, les ooïdes en demi-lune, les ooïdes "cassés" *sensu lato* ou les moules effondrés d'ooïdes (parfois appelés ooïdes "déformés").

#### Mots-clefs :

- ooïdes ;
- carbonates microbiens ;
- Salta ;
- Argentine ;
- Maastrichtien-Danien

<sup>3</sup> 93 avenue des acacias, 91800 Brunoy (France) lapointe-philippe@orange.fr



*Published online in final form (pdf) on January 20, 2024* [Editor: Michel Moullade; language editor: Phil Salvador; technical editor: Bruno R.C. GRANIER]

<sup>&</sup>lt;sup>1</sup> 2 impasse Charles Martel, 29217 Plougonvelin (France) brcgranier@free.fr

<sup>&</sup>lt;sup>2</sup> Dépt. STU, Fac. Sci. Tech., UBO, 6 avenue Le Gorgeu, CS 93837, F-29238 Brest (France) bgranier@univ-brest.fr



## **1. Introduction**

The petrographic analysis of thin sections reveals that there is still much to be gleaned from microbial carbonates (GRANIER & LAPOINTE, 2021, 2022a, 2022b); this is the core idea behind the KALKOWSKY Project. The material presented here documents a new occurrence of 'eccentric' ooids from the lacustrine (Maastrichtian to Danian) Yacoraite Formation (*e.g.*, CóNSOLE GONELLA *et al.*, 2012; FREIRE, 2012) in NW Argentina.

# 2. Material and general setting

The material under study was collected by one of us (P.L.) with the assistance of three IFP colleagues (Bernard COLLETTA, Jean LETOUZEY, and Roland VIALY) from two distinct localities in the provinces of Salta and Jujuy in NW Argentina (Fig. 1):

1) on October 6, 1988: The first section (Fig. 2), already documented in GRANIER and LAPOINTE (2022b, Figs. 3.D, 4-5), was found approximately 60 km south of Salta. It is situated on a bend of Road 47 from Coronel Moldes to Puente Dique Cabra Corral (Fig. 1.D), precisely at 25°17'04.4"S 65°24' 56.1"W (Province of Salta, Argen-

tina). This outcrop, located in the Metán subbasin of the Salta Basin, is referred to as "Afloramento Viñuales" of the "Sequência Balbuena IV" of the Yacoraite Formation (FREIRE, 2012, Figs. 5.1, 5.10, 8.7) and is assigned a Danian age. Four petrographic thin sections were prepared from two rock pieces labelled ARA 268 and ARA 269, collected near the top of the logged section. Although the first two thin sections (ARA 268 and ARA 269) are likely lost, two new thin sections (AG 268 and AG 269, registered as MHNG-GEPI-2024-10268 and 10269 in the collections of the Musée d'Histoire Naturelle de Genève, Switzerland) were prepared from offcuts of the initial two;

2) on October 15, 1988: The second section, measured by the same group of field geologists (Fig. 3.B), is exposed in a canyon located 6.5 km south of Palma Sola, west of the truck road connecting this locality to El Sauzal, approximately 100 km east of San Salvador de Jujuy (Fig. 1.C), at around 24°05'24.0"S 64°17'46.4"W (Province of Jujuy, Argentina). This second Yacoraite section is situated in the Lomas de Olmedo subbasin of the Salta Basin and is presumed to be of Maastrichtian age. Two petrographic thin sections were prepared from one rock piece labelled ARA 351, collected near the bottom of the exposed section



**Figure 1: A)** Location map of the provinces of Jujuy (red) and Salta (blue) in Argentina; **B)** location map of the sampling localities 268-269 in the Province of Salta and 351 in the Province of Jujuy; **C)** location of the sampling locality 351 in a canyon section 6.5 km south of Palma Sola, Province of Jujuy; **D)** location of the sampling localities 268-269 on a bend of Road 47 from Coronel Moldes to Puente Dique Cabra Corral, Province of Salta.

(Fig. 3.B-C). While the first thin section (ARA 351) is likely lost, a second thin section (AG 351, registered as MHNG-GEPI-2024-10351 in the collections of the Musée d'Histoire Naturelle de Genève, Switzerland) was prepared from an offcut of the original.

## 3. Descriptions of samples ARA 268, ARA 269, and ARA 351

Thin section ARA 268 (Fig. 4.A) reveals three stromatolitic microcolumns, each approximately 1 cm wide, containing silt and coated grains in the stromatolitic inner vugs and in the intercolumnar space. In contrast, the microfacies of thin sections AG 268, AG 269, and ARA 269 (Fig. 4.B) consist of 1) a floatstone of ooids and bothryoids (also spelled 'botryoids') with a silty matrix and 2) fibrous sparitic crusts growing on some bothryoids. The matrix also contains silt-sized quartz and some fish teeth.

The microfacies of both thin sections ARA 351 (Fig. 4.C) and AG 351 corresponds to a floatstone of bothryoids and oolitic lithoclasts with an oolitic grainstone matrix. Some lithoclasts exhibit a superficial oolitic coating. Ostracod shells are commonly observed as nuclei of ooids.

sions" following FOLK, 1974) cortical layers are thicker right above the nucleus, and thin laterally and downward. It appears that the center of mass of the ooid did not change with the addition of a new fibrite layer. The amber-yellow tint of the fibrite crystals is unquestionably related to organic content (GRANIER, 2020), with calcite fibers incorporating a diffuse organic network, possibly the remnants of Extracellular Polymeric Substances (EPS).

A bothryoid from AG 269 (Fig. 5.D) is composed of a cluster of ooids, including one asymmetric ooid with an off-center siliciclastic nucleus, showing similarities with the previous example. Initially, the latter likely formed a first aggregate with another ooid, subsequently forming a biooid (*cf.* GRANIER & LAPOINTE, 2022b). New ooids joined to form a larger aggregate, then a bothryoid.

Ooids and bothyroids with anisopachous fibrite cortical layers are common among the coated grains of AG 351 (Fig. 5.F). The cortex of another asymmetric ooid from ARA 351 (Figs. 5.G, 6.C) exhibits significant variation in the thickness of its outermost layers.

#### 5. Discussion

In this chapter, Argentinian (Jujuy and Salta) asymmetric ooids are discussed in terms of differences and similarities with some other specific ooid types: 'broken' ooids *sensu lato*, 'distorted' ooids, half-moon ooids, hiatus ooids, and wobbly ooids.

#### **Differences:**

1) 'Broken and regenerated' ooids (CAROZZI, 1961): As stated previously, the nuclei of two specimens from AG 268 (Fig. 5.B-C), which are found approximately 5 mm away in the same thin section (Fig. 5.A), represent the two halves of the same original ooid. Both pieces have undergone partial regeneration. However, whereas the first layer of the 'regenerated' cortex is continuous in typical 'broken and regenerated' ooids, the Salta specimens (Fig. 5.B-C) are characterized by the non-continuous nature of their outer cortical layers. The latter commonly abut against the inner layers and include noticeable gaps.

2) 'Distorted' ooids (CAYEUX, 1935) and 3) halfmoon ooids (WHERRY, 1915): Recently, GRANIER and coauthors (GRANIER *et al.*, 2022; GRANIER & KENDALL, 2022) demonstrated that some 'distorted ooids' are, in fact, collapsed oomolds, *i.e.*, a result of diagenetic processes involving leaching of the ooids followed by mechanical compaction. Similarly, half-moon ooids are formed through the leaching of oolitic cortices, causing the nuclei and some impurities to settle at the bottom of oomoldic cavities. Both types are associated to diagenetic processes. In contrast, the features observed in our Argentinian oolites are 'genetic', *i.e.*, indicating a relationship with synsedimentary growth processes.

**Figure 2:** Shematic drawing of the Salta section (Cabra Corral) with location of samples 269 and 270 in bold red (excerpt from GRANIER & LAPOINTE, 2022b).

dessiccation cracks

**4**269 (+ 270)

266 thin sandy layer

264 lime mudstone

lime mudstone

greenish claystones

stromatolites (at top)

and oolite

260 stromatolites and oolite

259 258 silty wackestone

257 oolite (grainstone)

green claystones

lime mudstone

thick stromatolitic layer and oolite

◀267 alternation of green claystones

alternation of lime mudstones and claystones (more or less silty)

lamellar stromatolites (at top)

(some load structures)

grey claystones or marlstones

with few thin limestone layers

and beige lime mudstone

272 distorted bedding (thin beds)

and marlstones

268

₹275

₹274

₹265

**4**273

**4**263

262

2/ 261

Ε

## 4. Descriptions of some ooids and bothryoids from thin sections AG 268, AG 269, AG 351, and ARA 351

The nucleus of one specimen from AG 268 (Figs. 5.B, 6.A) is a hemiooid *sensu* KALKOWSKY'S (1908) classification. Similarly to most 'broken' ooids *sensu stricto*, the break lines align with the calcite fibers of the cortical layers. In this case, a half-piece of the ooid has undergone partial regeneration, a phenomenon also observed in the second half (Fig. 5.C), which was found approximately 5 mm away in the same thin section (Fig. 5.A). However, both pieces distinctly differ from typical 'broken and regenerated' ooids, *i.e.*, 'broken' ooids *sensu lato*, due to the non-continuous nature of their 'regenerated' cortices.

The siliciclastic nucleus of another asymmetric ooid from AG 268 (Figs. 5.E, 6.B) is protruding. Yellowish 'fibrite' (a neologism for 'fibrous calcite' as coined by GRANIER and LAPOINTE, 2022a, *i.e.*, "material with one large and two small dimen-



4) Hiatus ooids: as defined by BERG (1944), such ooids exhibit some obliquely truncated cortical layers, suggesting that their asymmetry likely results from mechanical abrasion, indicative of erosional processes. Partly abraded layers of the inner cortex terminate beneath the boundary with the outer cortex. In contrast, in two ooids from AG 268 (Figs. 5.B-C, 6.A) some layers of the outer cortex terminate above the boundary with the inner cortex. More generally, the asymmetry observed in the Argentinian material is primarily associated with growth processes rather than abrasion.

dome-shaped stromatolites overhang in the cliff (?) limestone with a yellow patina (no description, out of reach) limestone with a pinkish patina fractures filled with secondary gypsum beef veins thin, scattered oolitic layers calcareous ooids mudstone limestone and ferruginous ooids (fair porosity) alternation of grey and pink beds, silty "competition" between oolites and stromatolites beige mudstone alternation of yellow mudstone and oolites dykes with grey muds injected downward limestone with large ooids undulaceous (sigmoidal) laminations in oolites yellow micrite irregular laminations in an oolitic limestone pack- to grainstone oolitic limestone (locally ooids have a cubic layout) bottom of the section = bed of the arroyo **∢**354 **∢**353 €352 Figure 3: The Jujuy section, a canyon section near Palma Sola: A) View of the canyon section; B)

5) Spiny ooids: According to DAVAUD and STRAS-SER (1990), "the external cortices are deformed and detached from the underlying cortices near the points of contacts between the grains", which "strongly suggest a postdepositional origin for the spines". This type of ooid is associated with early diagenetic processes.

stromatolites.

schematic drawing of the canyon section with location of sample 351

in red; C) lowermost oolites and





**Figure 4:** High resolution scans of the thin sections: **A)** three stromatolitic microcolumns, each approximately 1 cm wide, containing silt and coated grains in the stromatolitic inner vugs and in the intercolumnar space, ARA 268, **B)** floatstone of bothryoids and ooids with a silty matrix, ARA 269, **C)** floatstone of lumps and bothryoids with an oolitic grainstone matrix, ARA 351 (all likely lost). Scale bar for all scans = 5 mm.

6) Wobbly ooids: In the case of the Jujuy wobbly ooids (GRANIER & LAPOINTE, 2022a), the asymmetry is associated with the growth of micritic bumps, likely of microbial origin, and successive shifts of the center of gravity. In the material presented here, there are no micritic bumps; instead, incomplete yellowish 'fibrite' coatings are present. These coatings either thicken or thin and commonly abut against older layers. If these fibrite crusts were made of micrite, the Argentinian aymmetric ooids would unequivocally be classified as oncoids. In contrast to the previously described wobbly ooids (GRANIER & LAPOINTE, 2022a), the centers of mass in our Salta specimens of Figure 6.A-B, .D (samples AG 268 and AG 269) did not significantly move during the latest growth stages.

#### Similarities:

Salta asymmetric ooids (samples AG 268 and AG 269) exhibit some similarities with the modern "quiet water oölites from laguna Madre, Texas", as described by FREEMAN (1962). According to the latter, these asymmetric features "seem not to be the result of etching or abrasion but rather they appear to be primary features of these oölites" (*op. cit.*, p. 478). The specimen in figures 5.E and 6.B with its outlying siliciclastic nucleus (sample AG 268) shows even more striking similarities with certain ooids documented by FREEMAN (1962, Fig. 6, photomicrographs A and B). However, in contrast to FREEMAN's ooids, the cortices of which are composed of aragonite, the Argentinian coat-

ed grains were likely made of high-Mg calcite (GRA-NIER & LAPOINTE, 2022b).

It is worth mentioning that, whereas the nuclei of the ooids illustrated in figures 5.D-E and 6.B (samples AG 268 and AG 269) consist of siliciclastic grains, the ooid cortices never incorporated any silt-sized quartz grains, even when present in the matrix. This demonstrates that, unlike some stromatolites, ooids lack the capacity to agglutinate or bind such exogenous grains.

#### 6. Conclusion

The distinctive ooids from the Maastrichtian-Danian Yacoraite Formation in NW Argentina, as described here, belong to a unique class of 'eccentric' ooids. Unlike the wobbly ooids, the examples studied here do not exhibit any micritic bumps, and their fibrite cortical layers are not isopachous. Instead, a discontinuous, anisopachous fibrite coating and, eventually, an eccentric position for their center of gravity are determining factors to explain their cortical asymmetry. Because they should not be confused with 'broken' ooids sensu stricto, 'broken and regenerated' ooids (CAROZZI, 1961), i.e., 'broken' ooids sensu lato, 'distorted' ooids (CAYEUX, 1935), half-moon ooids (WHERRY, 1915), hiatus ooids (BERG, 1944), spiny ooids (Davaud & Strasser, 1990), or wobbly ooids (GRANIER & LAPOINTE, 2022a), it is recommended to simply categorize them as asymmetric ooids.







**Figure 5:** A-C, E) thin section AG 268: A) microfacies (the two half ooids are arrowed); B-C; broken and asymmetrically regenerated ooids; E) asymmetric ooid with a protruding siliciclastic nucleus (arrowed); D) thin section AG 269: bothryoid composed of a cluster of ooids, including one asymmetric ooid with its off-center siliciclastic nucleus (arrowed); F) thin section AG 351: ooid with an asymmetric cortex at the center of the photomicrograph; G) thin section ARA 351: ooid with an asymmetric, non-continuous cortex. A-E: Road 47 from Coronel Moldes to Puente Dique Cabra Corral, Province of Salta; F-G: south of Palma Sola, Province of Jujuy. A) scale bar = 1 mm; B-F) scale bar = 250  $\mu$ m; G) scale bar = 100  $\mu$ m.

### Acknowledgements

The rock samples studied here were collected by the second author (P.L.) on the occasion of a joint mission of Total - Compagnie Française des Pétroles, and IFP - Institut Français du Pétrole from October 5 to November 3, 1988. He acknowledges the support of his IFP colleagues, Bernard COLLETTA, Jean LETOUZEY, and Roland VIA- LY, for fieldwork. The authors acknowledge the detailed reviews provided by Christopher G.St.C. KEN-DALL (University of South Carolina) and André STRAS-SER (Université de Fribourg/Univesität Freiburg) and their help in improving this short paper. The first author (B.G.) would like to thank Phil SALVA-DOR for his appreciated help with the final (English) text.



**Figure 6: A-B)** thin section AG 268, Road 47 from Coronel Moldes to Puente Dique Cabra Corral, Province of Salta: A) following the ooid breakage, growth of the regenerated cortex is restricted to a part of the fracture plane, one edge and the convex part, see Fig. 5.B; B) the center of mass of the ooid is likely the protruding siliciclastic nucleus controling the upward growth of the cortex, see Fig. 5.E; C) cauliflower-like developments on the outermost cortical layers, thin section ARA 351, south of Palma Sola, Province of Jujuy, see Fig. 5.G. Scale bar for all photomicrographs = 500  $\mu$ m.

#### **Bibliographic references**

- BERG G. (1944).- Vergleichende Petrographie oolithischer Eisenerze.- *Archiv für Lagerstättenforschung*, Berlin, Heft 76, p. 7-128 (6 Pls.).
- CAROZZI A.V. (1961).- Oolithes remaniées, brisées et régénérées dans le Mississippien des chaînes frontales, Alberta Central, Canada.- Archives des Sciences, Genève, vol. 14, no. 2, p. 281-296.
- CAYEUX L. (1935).- Les roches sédimentaires de France. Roches carbonatées.- Masson & Cie, Paris, 447 p.
- Cónsole Gonella C.A., GRIFFIN M., CIONE A., GOUI-RIC CAVALLI S. & ACEÑOLAZA F.G. (2012).- Paleontología de la Formación Yacoraite (Maastrichtiano-Daniano) en el ámbito de la Subcuenca de Tres Cruces, Cordillera Oriental de la provincia de Jujuy, Argentina.- Relatorio de la XIII Reunión Argentina de Sedimentología (16-19 de mayo de 2012), Salta, p. 45-56.
- DAVAUD R.L. & STRASSER A. (1990).- Spiny ooids: Early subaerial deformation as opposed to late burial compaction.- *Geology*, Boulder - CO, vol. 18, no. 9, p. 816-819.



- FOLK R.L. (1974).- The natural history of crystalline calcium carbonate: Effect of magnesium content and salinity.- *Journal of Sedimentary Petrology*, Tulsa - OK, vol. 44, no. 1, p. 40-53.
- FREIRE E.B. (2012, unpublished).- Caracterização estratigráfica em alta resolução das sequências calcárias de origem microbiana do intervalo paleocênico da Formação Yacoraite (Sequência Balbuena IV) na região de Salta - Argentina.-MSc Dissertação, UFRJ, Rio de Janeiro, 243 p.
- FREEMAN T. (1962).- Quiet water oolites from Laguna Madre, Texas.- Journal of Sedimentary Petrology, Tulsa - OK, vol. 32, no. 3, p. 475-483.
- GASIEWICZ A. (1984a).- Górnojurajskie ooidy o niewspółśrodkowych powłokach.- *Kwartalnik Geologiczny*, Biłgoraj, t. 28, no. 1, p. 93-106.
- GASIEWICZ A. (1984b).- Eccentric ooids.- Neues Jahrbuch für Geologie und Paläontologie, Monatshefte, Stuttgart, Heft 4, p. 204-211.
- GRANIER B. (2020).- The biosignature of sparite permits the distinction between gravitational cement and endostromatolites.- *Carnets Geol.*, Madrid, vol. 20, no. 20, p. 407-419. DOI: 10.2110/carnets.2020.2020
- GRANIER B.R.C., KENDALL Ch.G.St.C. & VACHARD D. (2022).- The KALKOWSKY Project - Chapter IV. Case study of the Happy Spraberry oil reservoirs in NW Texas (with a micropaleontologic and biostratigraphic supplement): Collapsed molds should not be treated as a category of distorted ooids.- *Carnets Geol.*, Madrid, vol. 22, no. 8, p. 311-332. DOI: 10.2110/carnets. 2022.2208
- GRANIER B.R.C. & KENDALL Ch.G.St.C. (2022).- Discussion of LAYA et al. (2021), Dissolution of ooids in seawater-derived fluids an example from Lower Permian re-sedimented carbonates, West Texas, USA [Sedimentology 68(6), 2671-2706].- Carnets Geol., Madrid, vol. 22, no. 10, p. 395-408. DOI: 10.2110/carnets. 2022.2210

- GRANIER B.R.C. & LAPOINTE Ph. (2021).- The KAL-KOWSKY Project - Chapter I. Ooid - stromatoid relationship in a stromatolite from the Maiz Gordo Fm (Argentina).- *Carnets Geol.*, Madrid, vol. 21, no. 9, p. 193-201. DOI: 10.2110/ carnets.2021.2109
- GRANIER B.R.C. & LAPOINTE Ph. (2022a).- The KAL-KOWSKY Project - Chapter II. Wobbly ooids in a stromatolite from the Yacoraite Formation (Argentina).- *Carnets Geol.*, Madrid, vol. 22, no. 3, p. 111-117. DOI: 10.2110/carnets.2022.2203
- GRANIER B.R.C. & LAPOINTE Ph. (2022b).- The KAL-KOWSKY Project - Chapter III. Significance of primary radial fabrics associated with ancient partly leached or recrystallized calcareous ooids.- *Carnets Geol.*, Madrid, vol. 22, no. 5, p. 149-160. DOI: 10.2110/carnets.2022.2205
- KALKOWSKY E. (1908).- 3. Oolith und stromatolith im norddeutschen Buntsandstein.- Zeitschrift der deutschen geologischen Gesellschaft, Berlin, Band 60, Heft I, p. 68-125 (Pls. IV-XI). URL: https://archive.org/details/zeitschriftderd 601908deut
- MARQUILLAS R.A., PAPA C. del & SABINO I.F. (2005).-Sedimentary aspects and paleoenvironmental evolution of a rift basin: Salta Group (Cretaceous-Paleogene), northwestern Argentina.-*International Journal of Earth Sciences (Geologische Rundschau*), vol. 94, p. 94-113.
- MORENO J. (1970).- Estratigrafía y paleogeografía del Cretácico Superior en la cuenca del norte argentino, con especial mención de los Subgrupos Balbuena y Sánta Barbara.- Revista de la Asociación Geológica Argentina, Buenos Aires, vol. 25, no. 1, p. 9-44.
- WHERRY E.T. (1915).- A peculiar oolite from Bethlehem, Pennsylvania.- Proceedings of The United States National Museum, Washington -DC, vol. 49 (1916), p. 153-156 (Pls. 40-41). URL: https://www.biodiversitylibrary.org/page/ 15692240